

Information theory application to inversions of acoustic data from a continental shelf environment

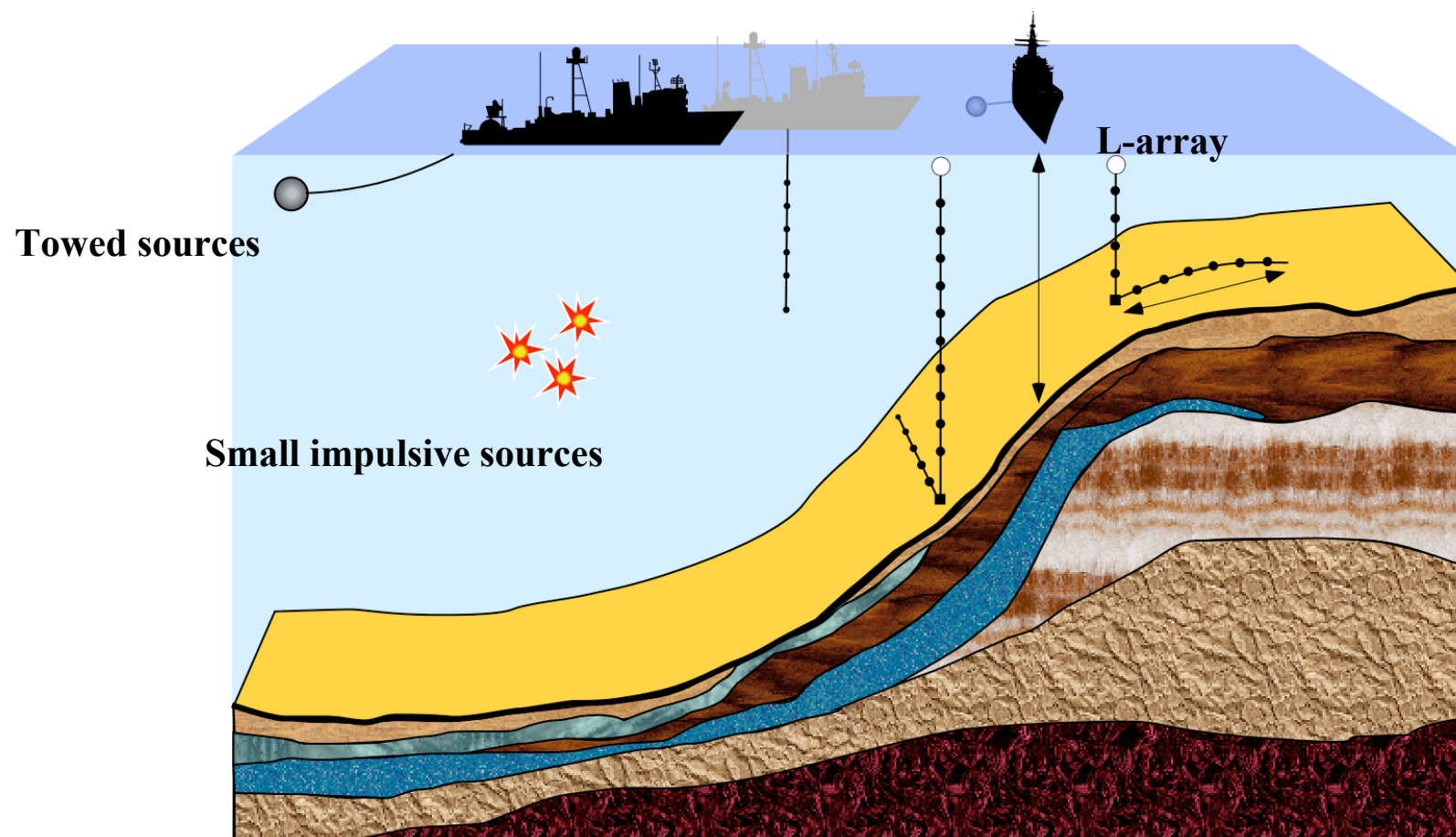
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- Inhomogeneous Oceans: Measurements and Inference
- Maximum entropy versus Bayesian
- Initial Computations
- Example maximum likelihood range-dependent calculations
- Summary

Example of acoustic measurements to infer parameters for inhomogeneous seabed



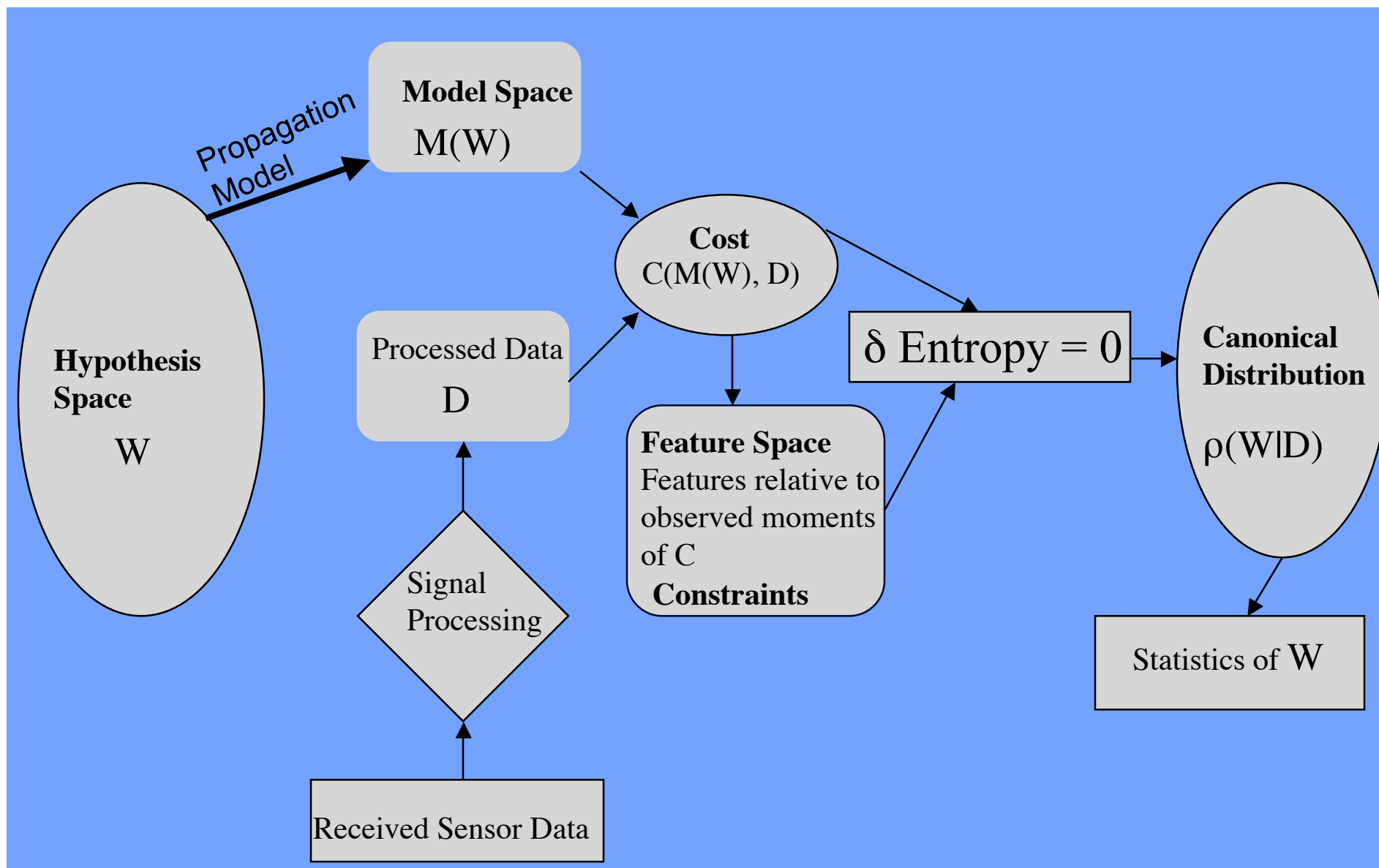
Measurements and Inferences

Why is inversion by itself not sufficient?

- Scientific objective: Interpret acoustic propagation in inhomogeneous ocean waveguides
 - Mode coupling mechanisms on shelf
 - Kramers Kronig relationships in seabed acoustics
- A maximum likelihood or inversion solution for C_{\min} is useful but insufficient
- **Uncertainty** in inferences is a natural consequence of
 - Environmental variability
 - Noise in data
 - Source-receiver motion and source level variability
 - Model errors
- Thus the need for posterior probability $\rho(W|D)$ that given data D the correct solution is W

- Bayesian approach designed to solve this problem
 - Requires likelihood function
- Alternative method is maximum entropy principle
 - Requires constraints

Maximum Entropy Approach to Uncertainty in Ocean Acoustics



A Canonical Distribution Approach (1)

Claude Shannon



Shannon or Gibbs Entropy

$$\delta S = 0 \text{ subject to stated constraints}$$

Edwin Jaynes



$$S = - \int_{\Omega} dW \rho(W|D) \ln [\rho(W|D) / \rho(W)]$$

Constraints

$$\int_{\Omega} dW \rho(W|D) = 1$$

$$\int_{\Omega} dW C(W) \rho(W|D) = \langle C \rangle = \frac{1}{2} (C_{min} + \bar{C})$$

C_{min} is global minimum determined from simulated annealing

\bar{C} average value of cost function space = $1/N \sum C(W_i)$

Analogy with statistical mechanics
for a closed system in thermodynamic
equilibrium with heat reservoir

$\delta S = 0$ subject to stated constraints

$$\rho(W|D) = \frac{\rho(W) \exp(-C(W, D)/T)}{Z}$$

Canonical Distribution

$$Z = \int_{\Omega} dW \rho(W) \exp(-C(W, D)/T)$$

Average $\langle C \rangle$ constraint determines T

$$\int_{\Omega} dW C(W) \rho(W|D) = \langle C \rangle$$

Relationship to Bayes formula

Thomas Bayes



$$\rho(W|D) = \rho(W) \rho(D|W) / \rho(D)$$

$$\rho(D) = \int_{\Omega} dW \rho(W) \rho(D|W)$$

$$\rho(W|D) = \rho(W) \rho(D|W) \left\{ \int_{\Omega} dW \rho(W) \rho(D|W) \right\}^{-1}$$

$$\rho(D|W) = \exp(-C/T)$$

Canonical distribution (not normalized) *plays role* of likelihood function

Average, standard deviation, and marginal distributions

$$\langle X \rangle = \frac{\int_{\Omega} d\mathbf{W} X(\mathbf{W}) \exp[-C(\mathbf{M}(\mathbf{W}), \mathbf{D})/T]}{Z}$$

$$\sigma_X = \frac{\int_{\Omega} d\mathbf{W} [X(\mathbf{W}) - \langle X \rangle]^2 \exp[-C(\mathbf{M}(\mathbf{W}), \mathbf{D})/T]}{Z}$$

Continuous formulation

$$P(w_i) = \frac{\int_{\Omega} d\mathbf{W}' \delta(w'_i - w_i) (\exp[-C(\mathbf{W}')/T])}{Z}$$

$$\langle X \rangle = \frac{\sum_i^N X(\mathbf{W}_i) \exp[-C(\mathbf{W}_i)/T]}{\sum_k^N \exp[-C(\mathbf{W}_k)/T]}$$

$$\sigma_X^2 = \frac{\sum_i^N [X(\mathbf{W}_i) - \langle X \rangle]^2 \exp[-C(\mathbf{W}_i)/T]}{\sum_k^N \exp[-C(\mathbf{W}_k)/T]}$$

Monte-Carlo integration

$$P(w_i) = \frac{\sum_{j \in_j w_i = w_i}^N \exp[-C(\mathbf{W}_j)/T]}{\sum_k^N \exp[-C(\mathbf{W}_k)/T]}$$

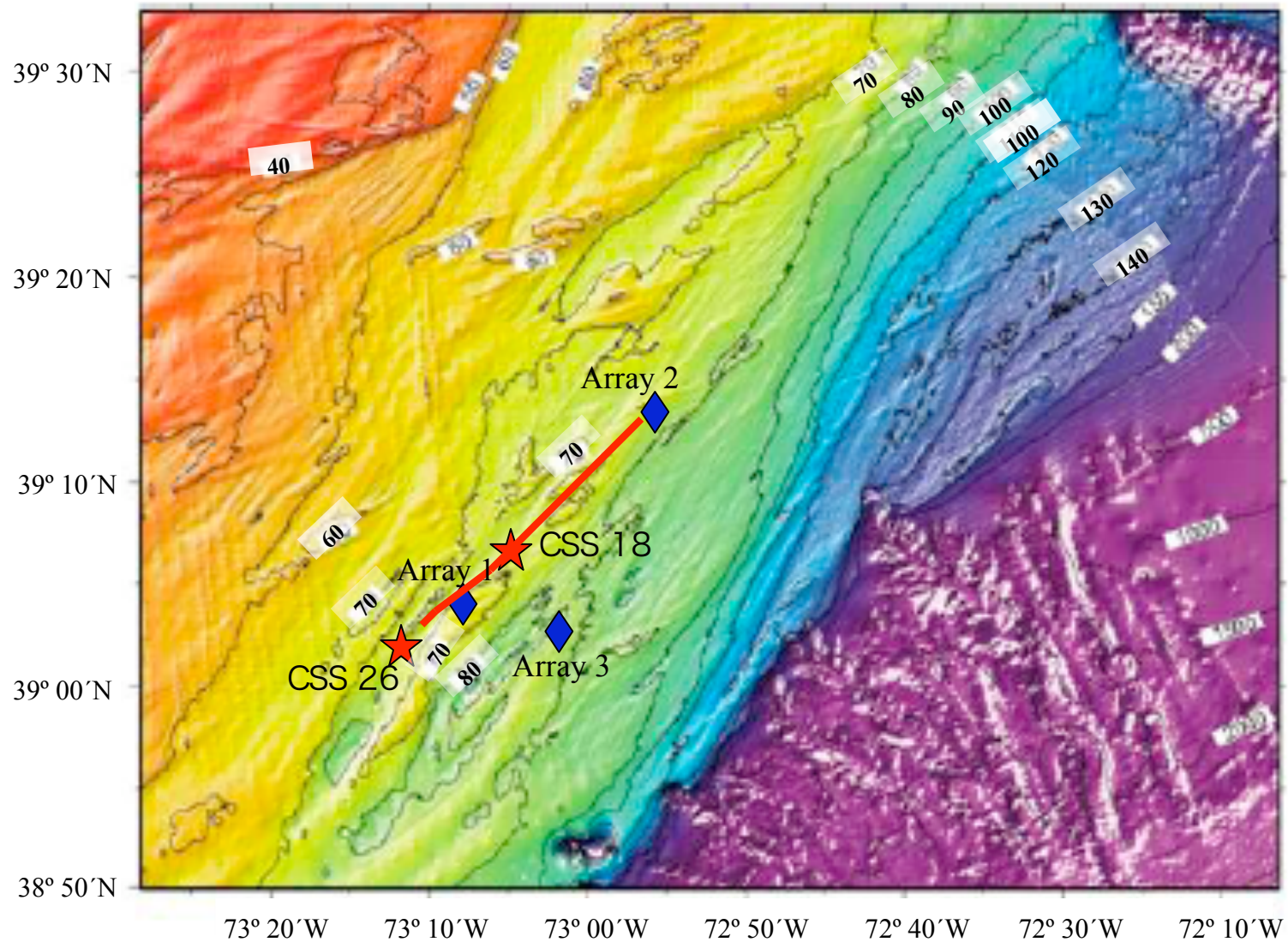
Pros and cons of Bayesian versus canonical distribution approach

- When prior information on noise and model errors is available, Bayesian approach is well justified
- Maximum entropy method appears well suited for problems with sparse data
 - Does not require direct assumptions about model / data errors or noise
 - Indirectly includes such information via constraints from *observed* features of cost
 - Prior information on $p(W)$ is included naturally via relative Shannon entropy
 - Leads to most conservative distribution
 - No restrictions on cost functions
 - Posterior distribution depends on cost function
 - Can include higher order moments of features, if available, via constraints

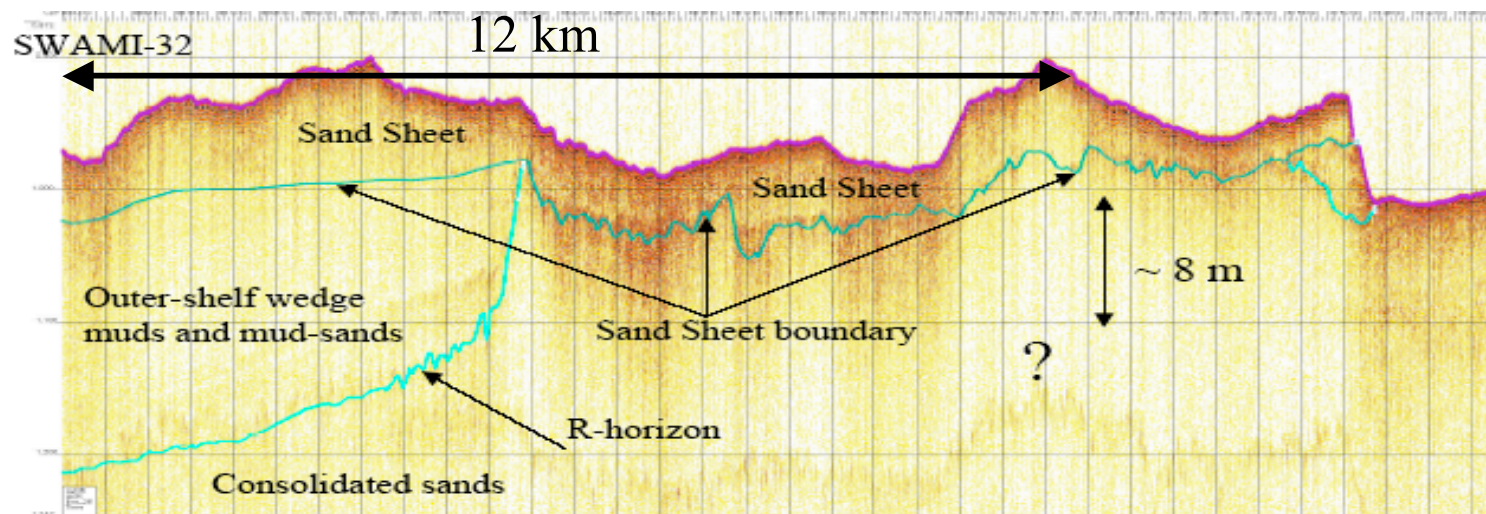
Application to acoustic data taken on continental shelf

- Shallow Water 2006 experiment
 - Data set of interest because of large spatial and temporal inhomogenities on continental shelf
 - Seabed
 - Water column
- Current Work
 - Maximum entropy principle applied to data with small range inhomogenities assuming range independence
 - Cross-slope range-dependent data using knowledge gained from MEP analysis, Goff geophysics characterization, and SSP measurements
 - Representation of range-dependent media - balance of representation versus number of parameters
 - Working to implement faster propagation model than PE

Location of sources and receivers

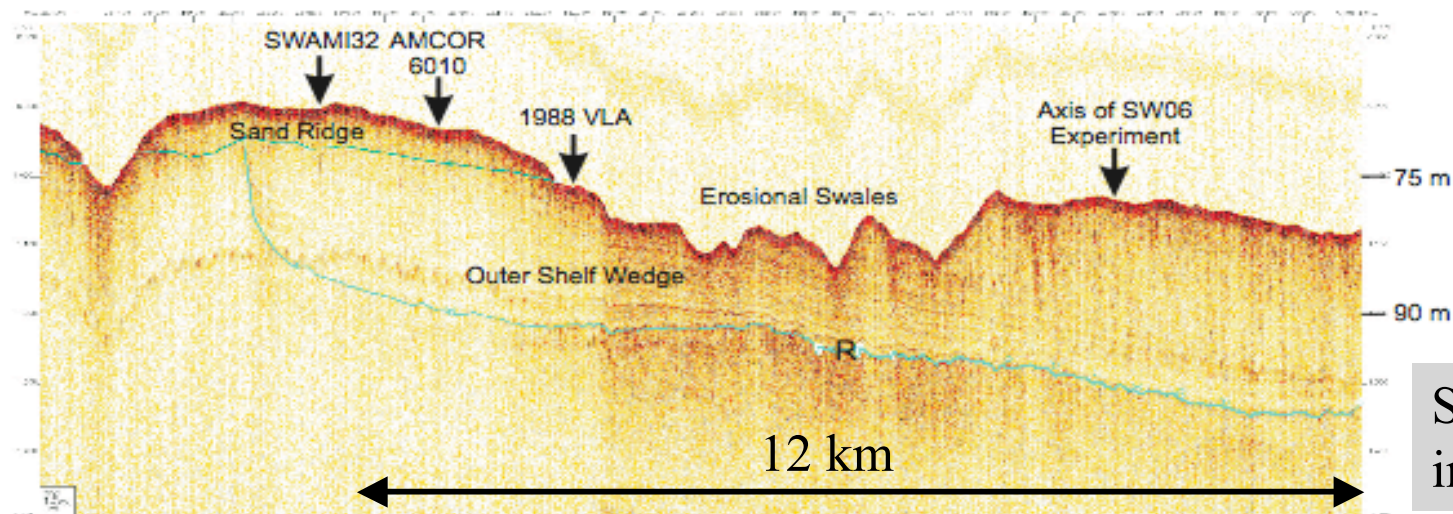


Chirp seismic reflection profiles



Track 1

Weak variations in SSP along track

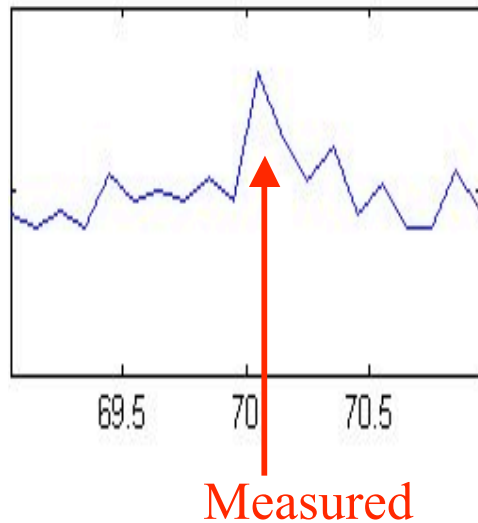


Track 2

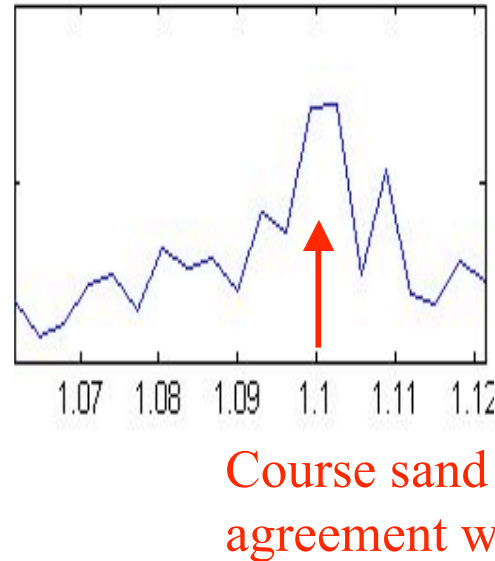
Strong variations in SSP along track

Marginal distributions from MEP for short range data (1-4 km) taken on Array 2

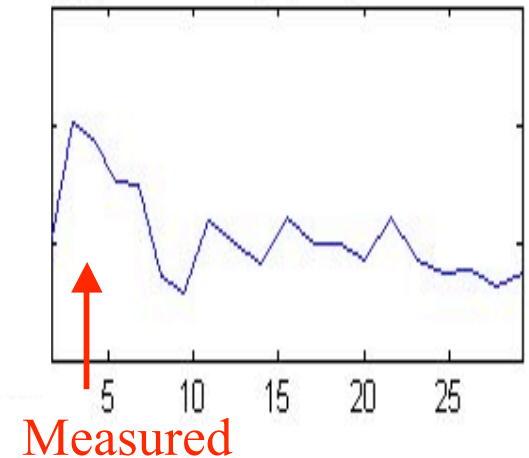
Water depth - m



Ratio(layer 1)



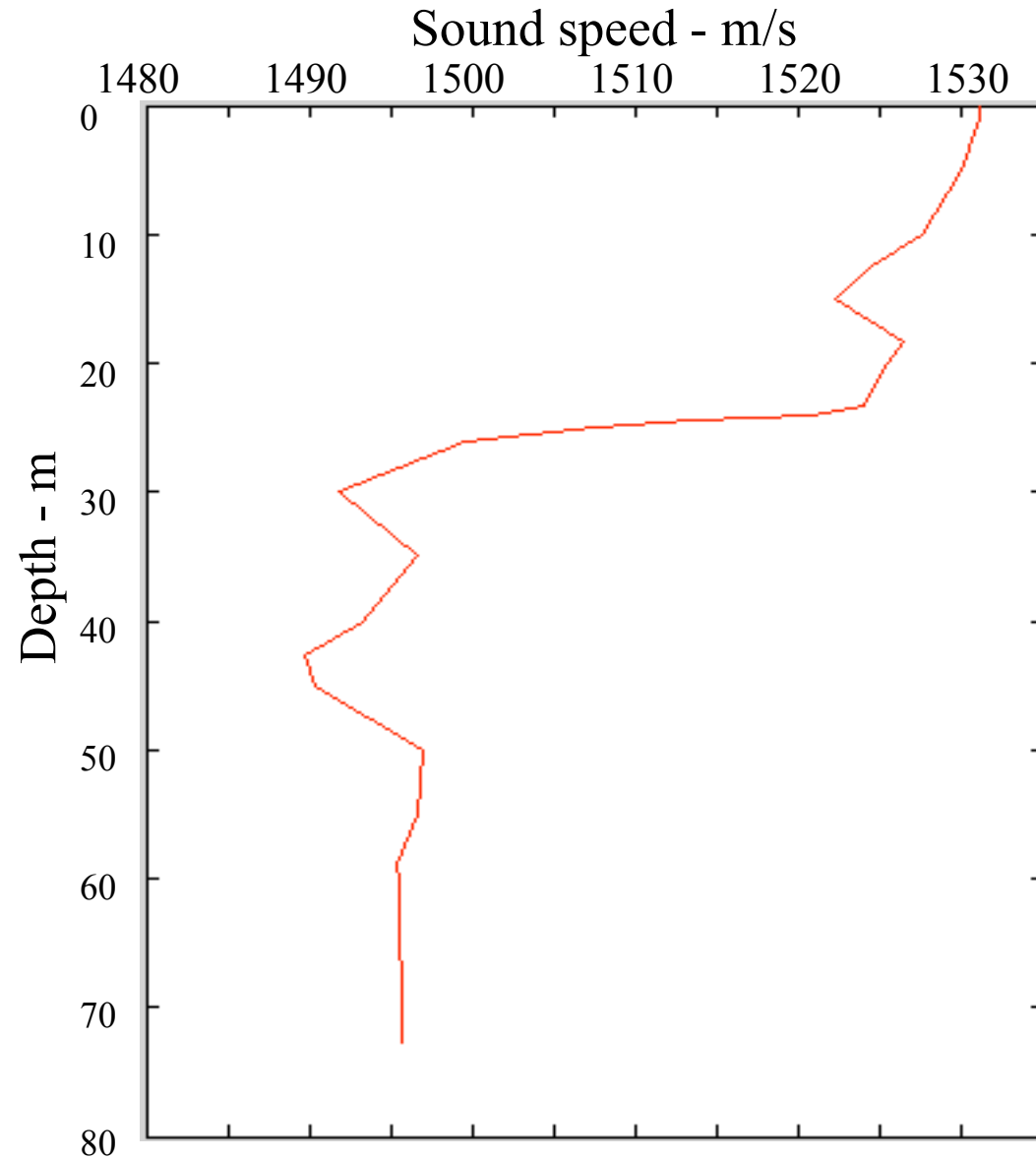
Thickness(layer1) - m



$$P(w_i) = \frac{\sum_{j \in \mathcal{J} : w_i = w_j} \exp[-C(\mathbf{W}_j)/T]}{\sum_k^N \exp[-C(\mathbf{W}_k)/T]}$$

2x10⁶ samples

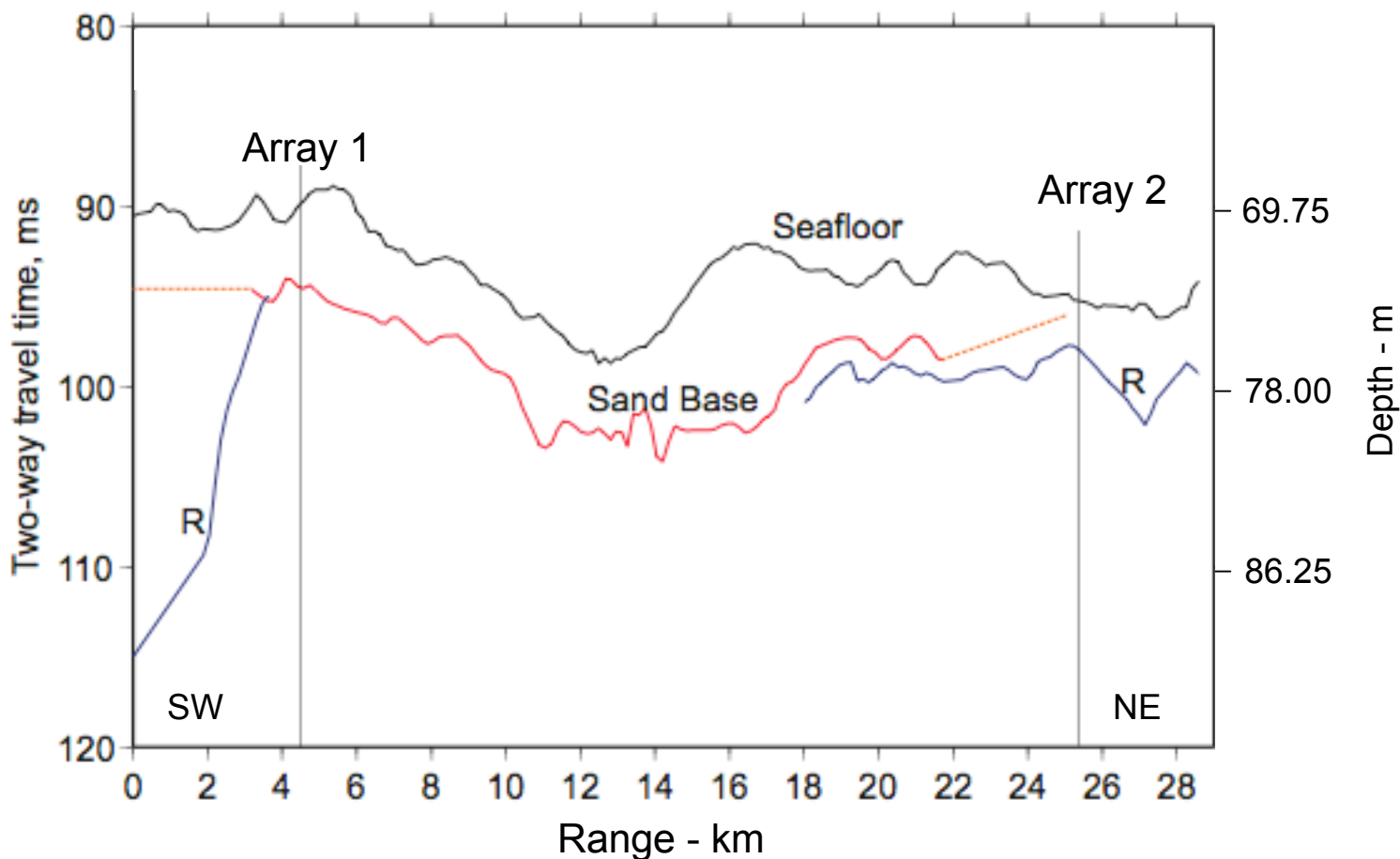
Sound speed profile along Track 1



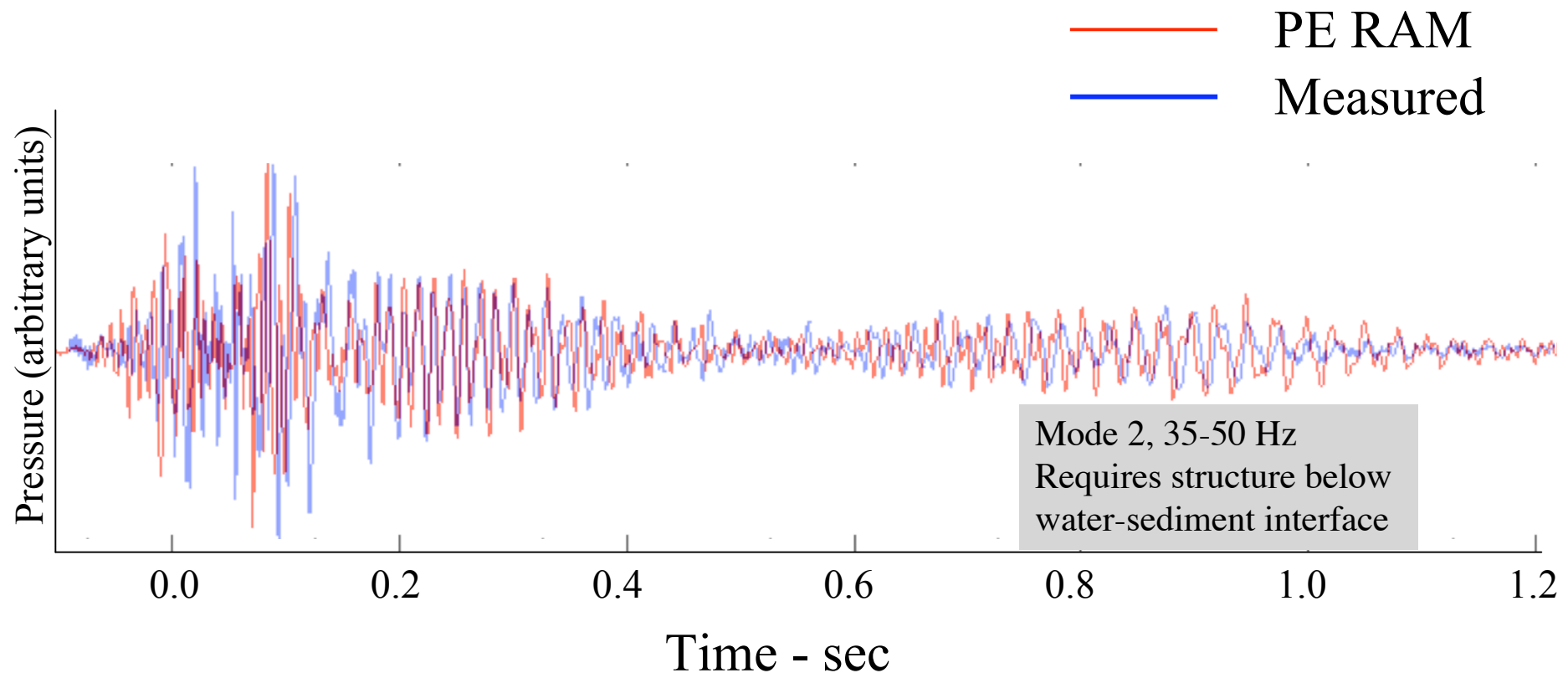
SSP had small variations along isobath

Geophysical structure of propagation track 1 from chirp reflection sonar

Weakly range-dependent track

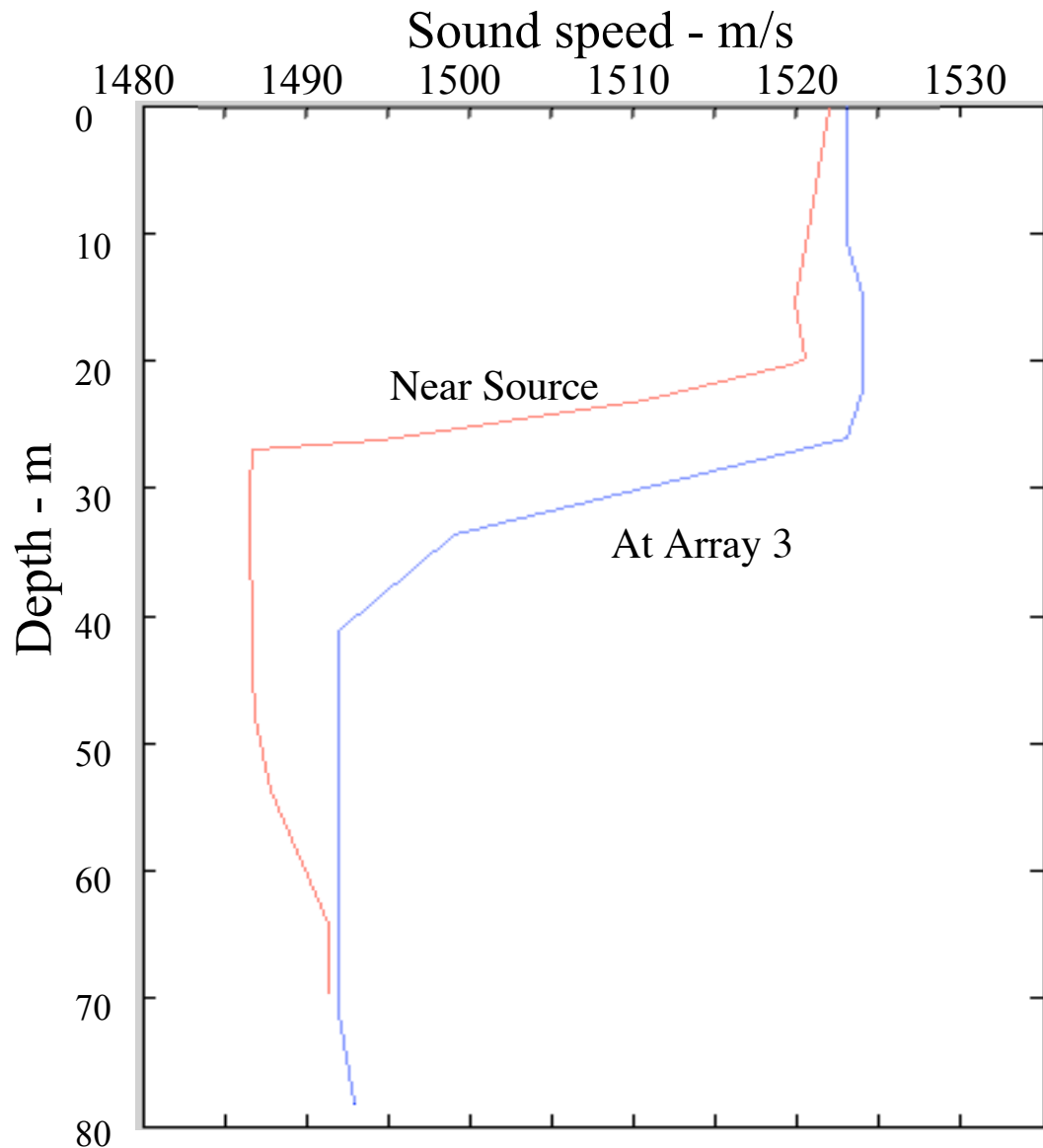


CSS event 26 model-data comparison for weakly range dependent track range 26.3 km, 35-325 Hz band



Information from short-range data with additional information provided by direct geophysical and sound speed measurements is sufficient for acoustic prediction along ~ isobath

Sound speed profile along Track 3

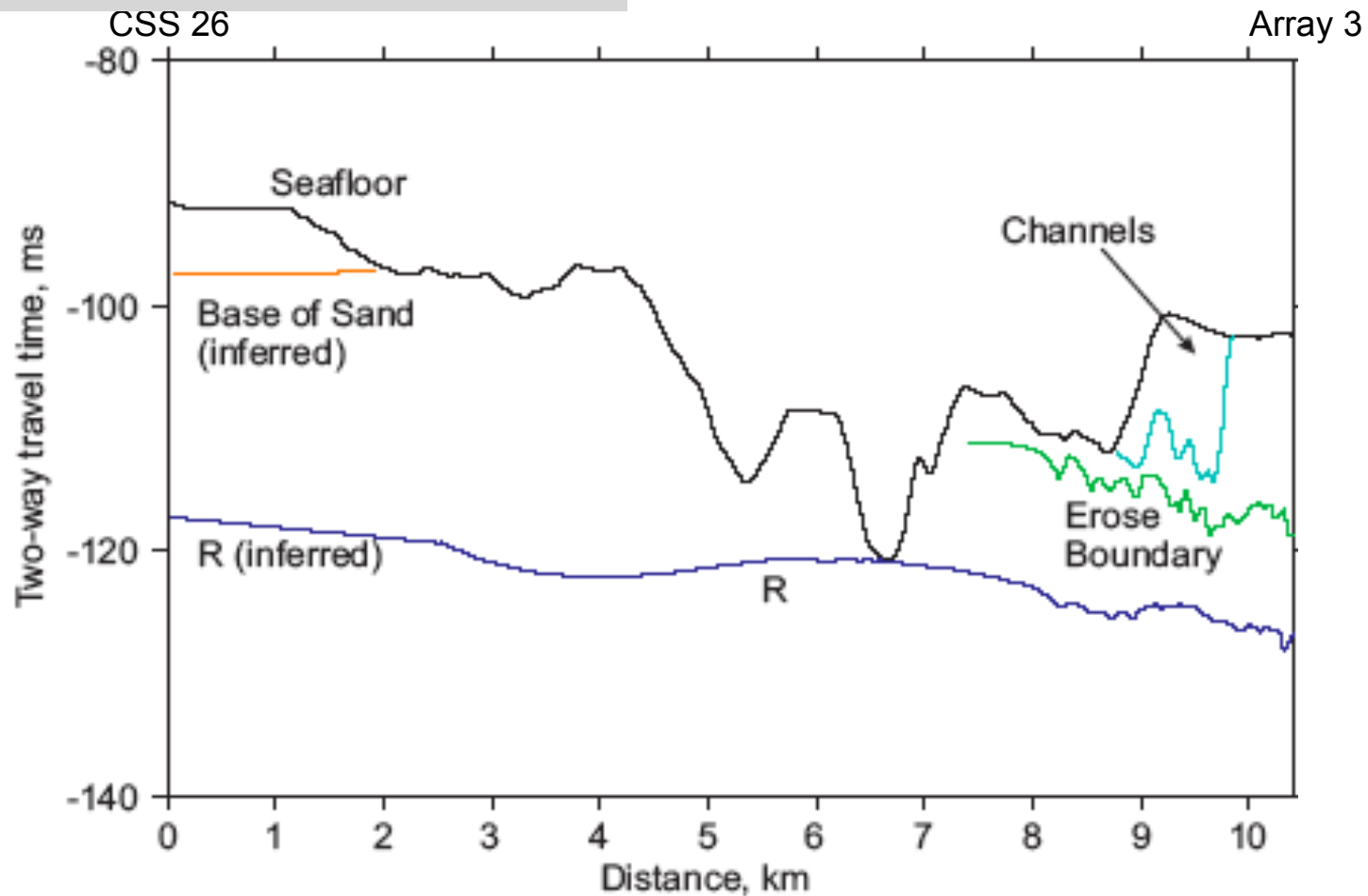


Stronger SSP
variation across
shelf. Similar SSPs -
but displacement
of thermocline

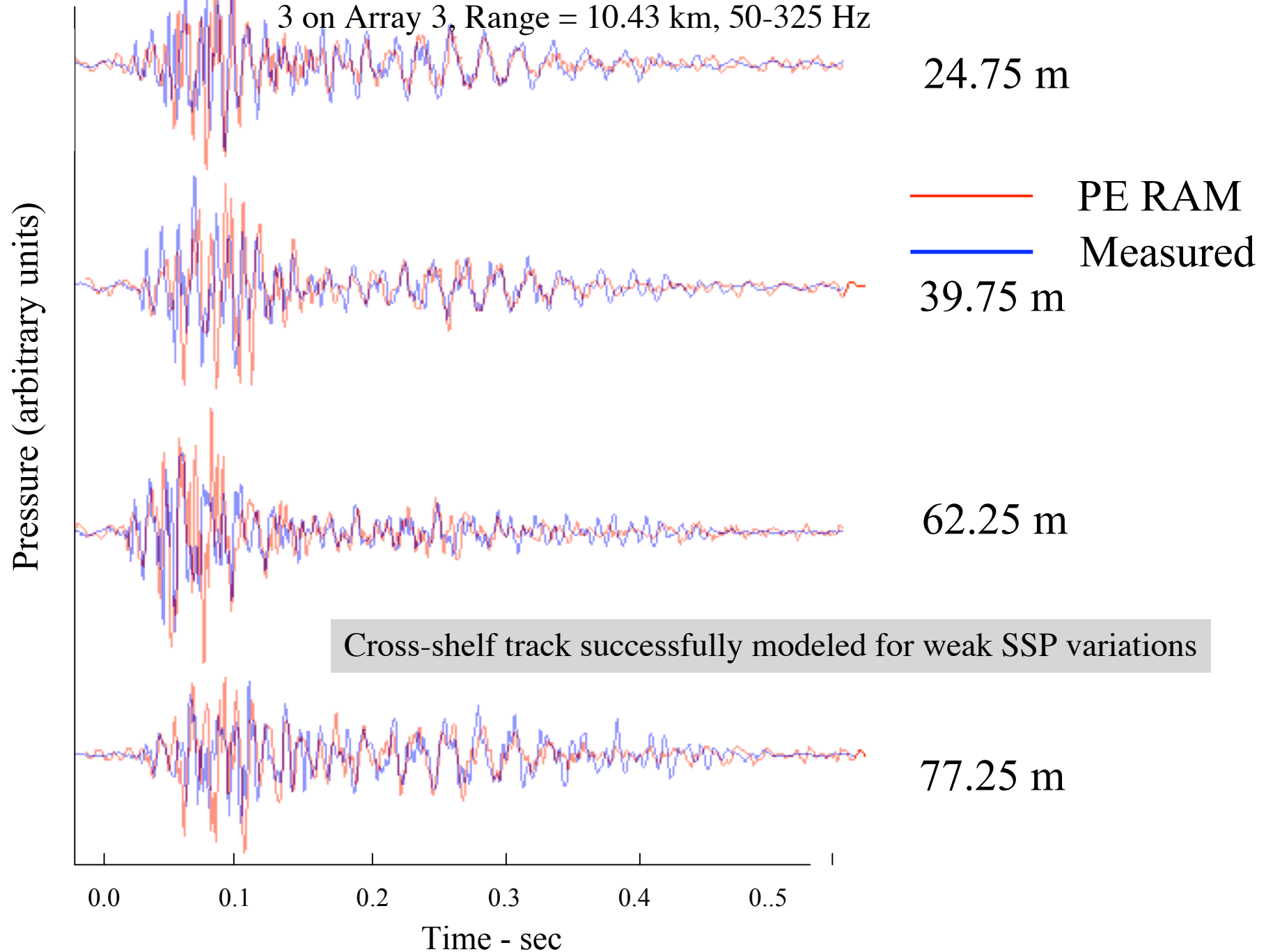
Hypothesis:
Place measured SSPs
at source and receiver and
interpolate for points between
source and receiver

Geophysical structure of propagation track 3 from chirp reflection sonar

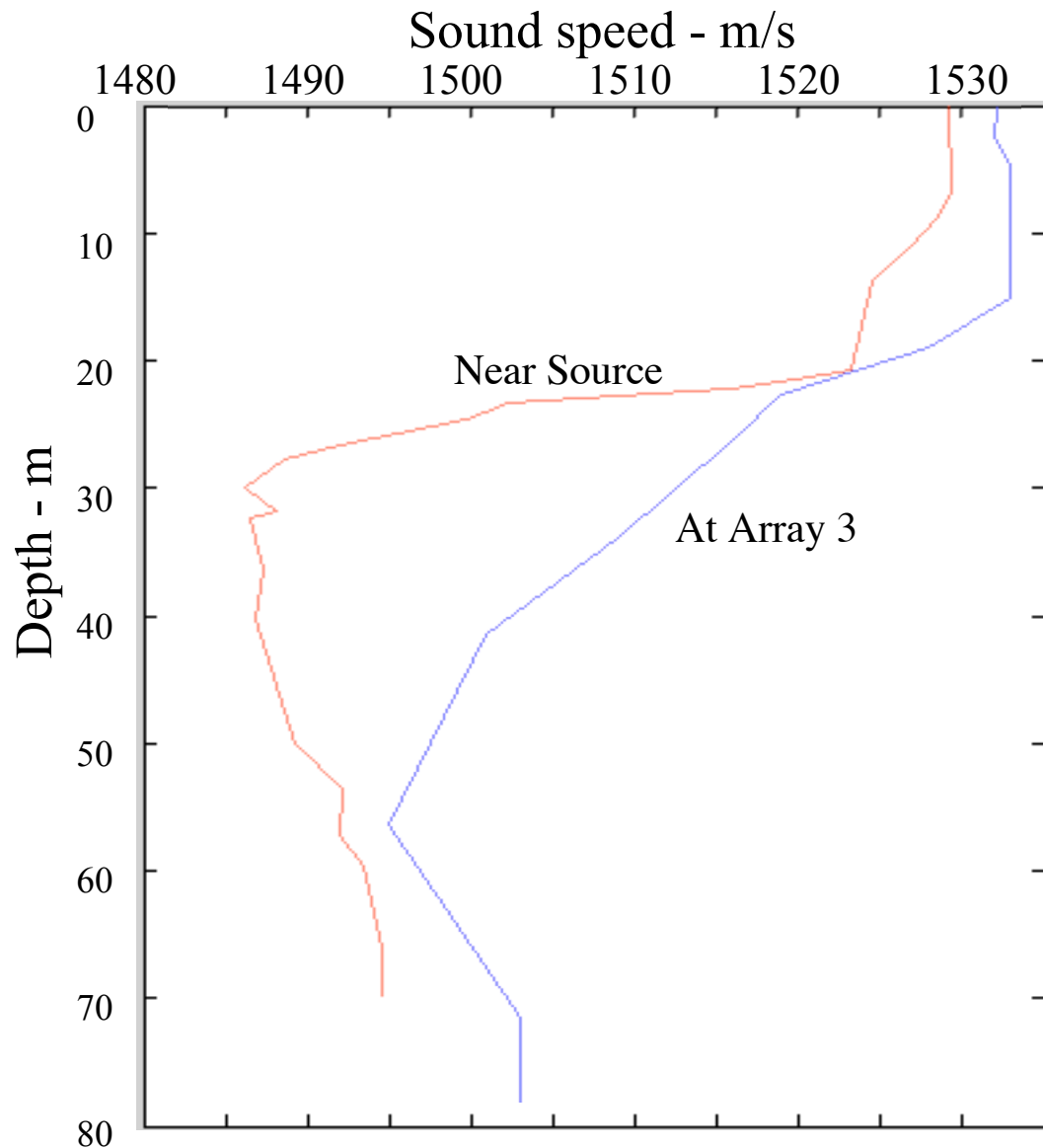
Moderately range-dependent track



Model Data Comparison of Received Time Series from CSS 26 on Track
3 on Array 3, Range = 10.43 km, 50-325 Hz



Sound speed profile along Track 2

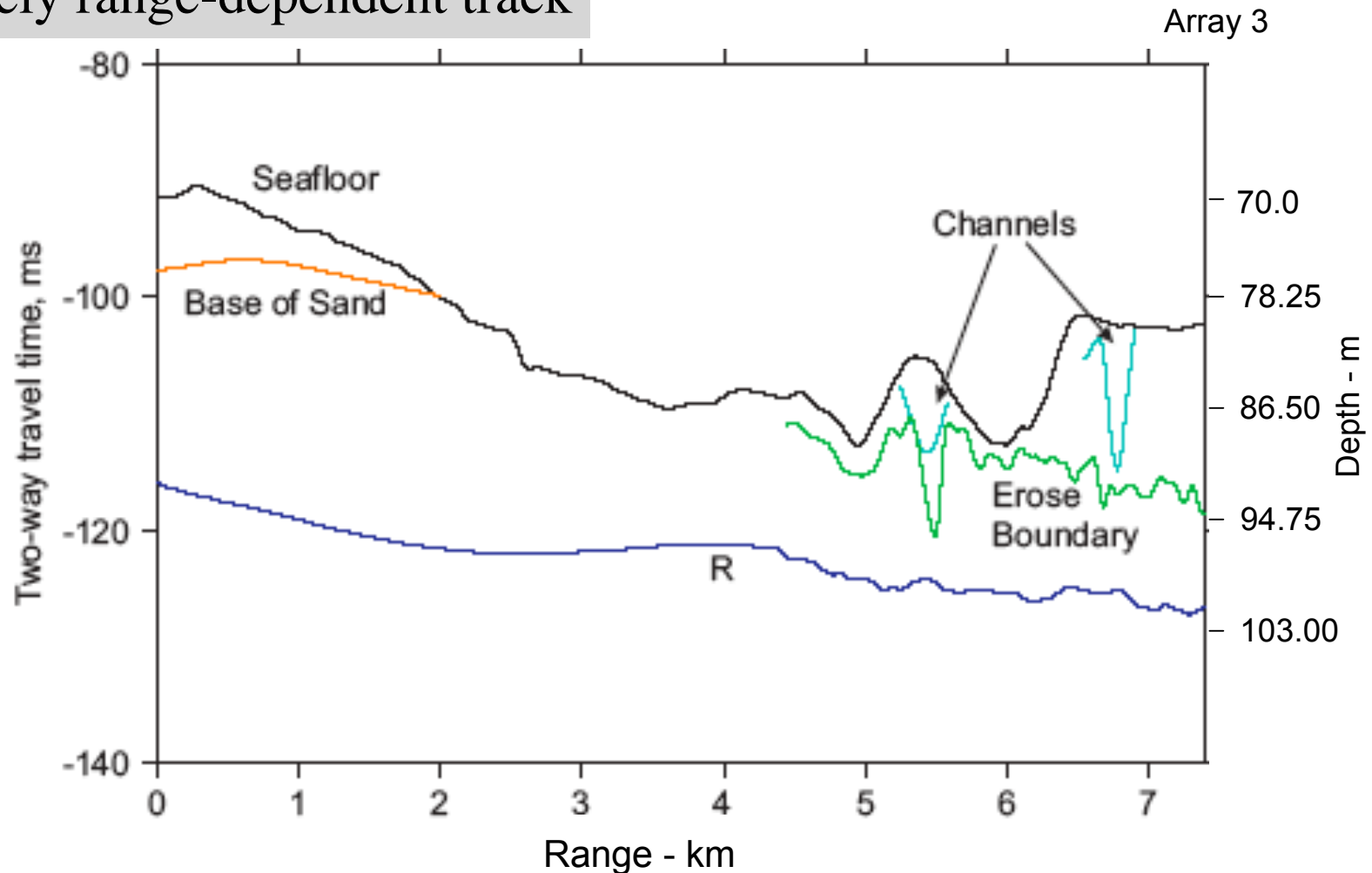


Strong SSP
variation across
shelf for track 2

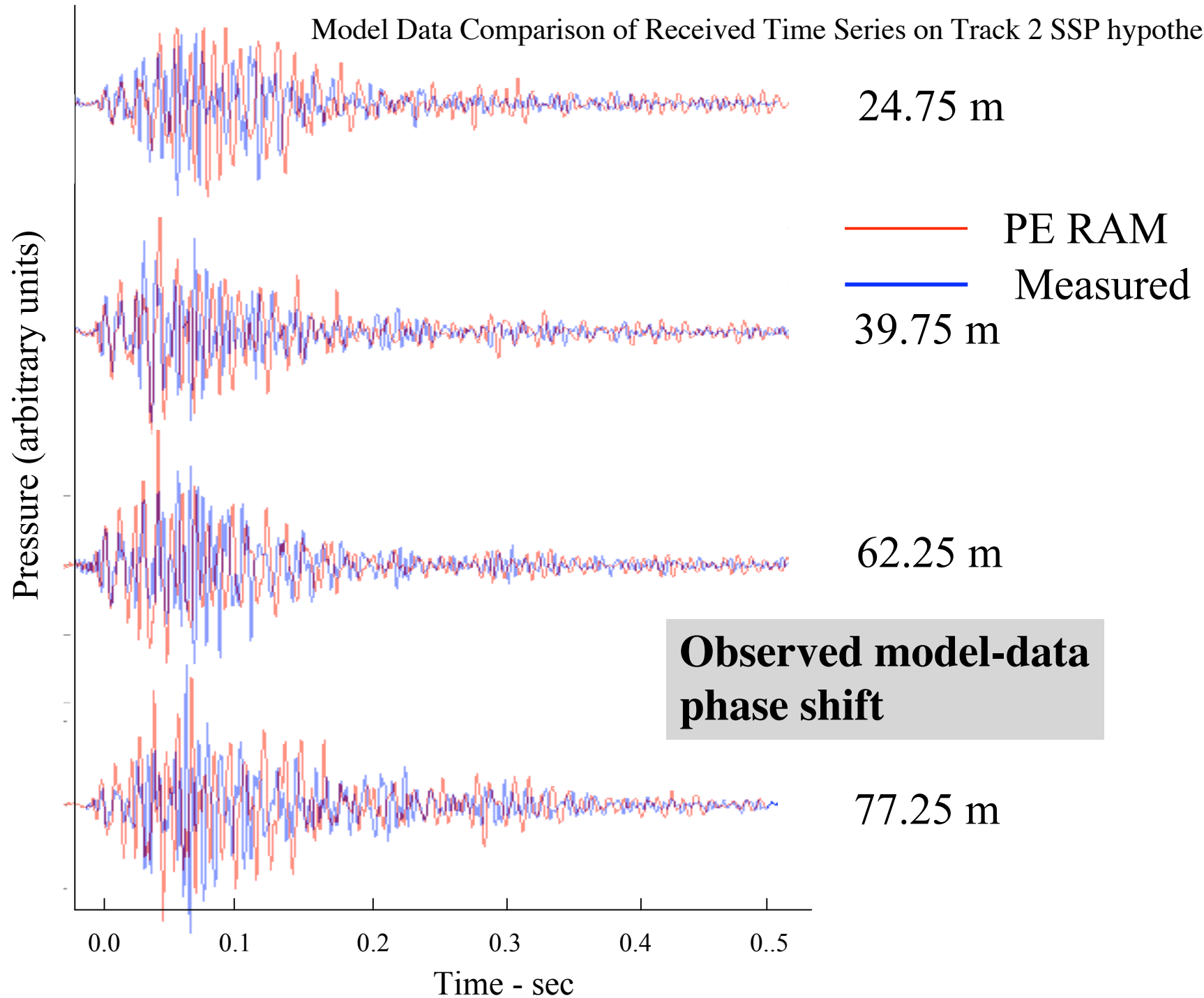
Hypothesis 1:
Place measured SSP
at source and receiver and
interpolate for points between
source and receiver

Geophysical structure of propagation track 2 from chirp reflection sonar

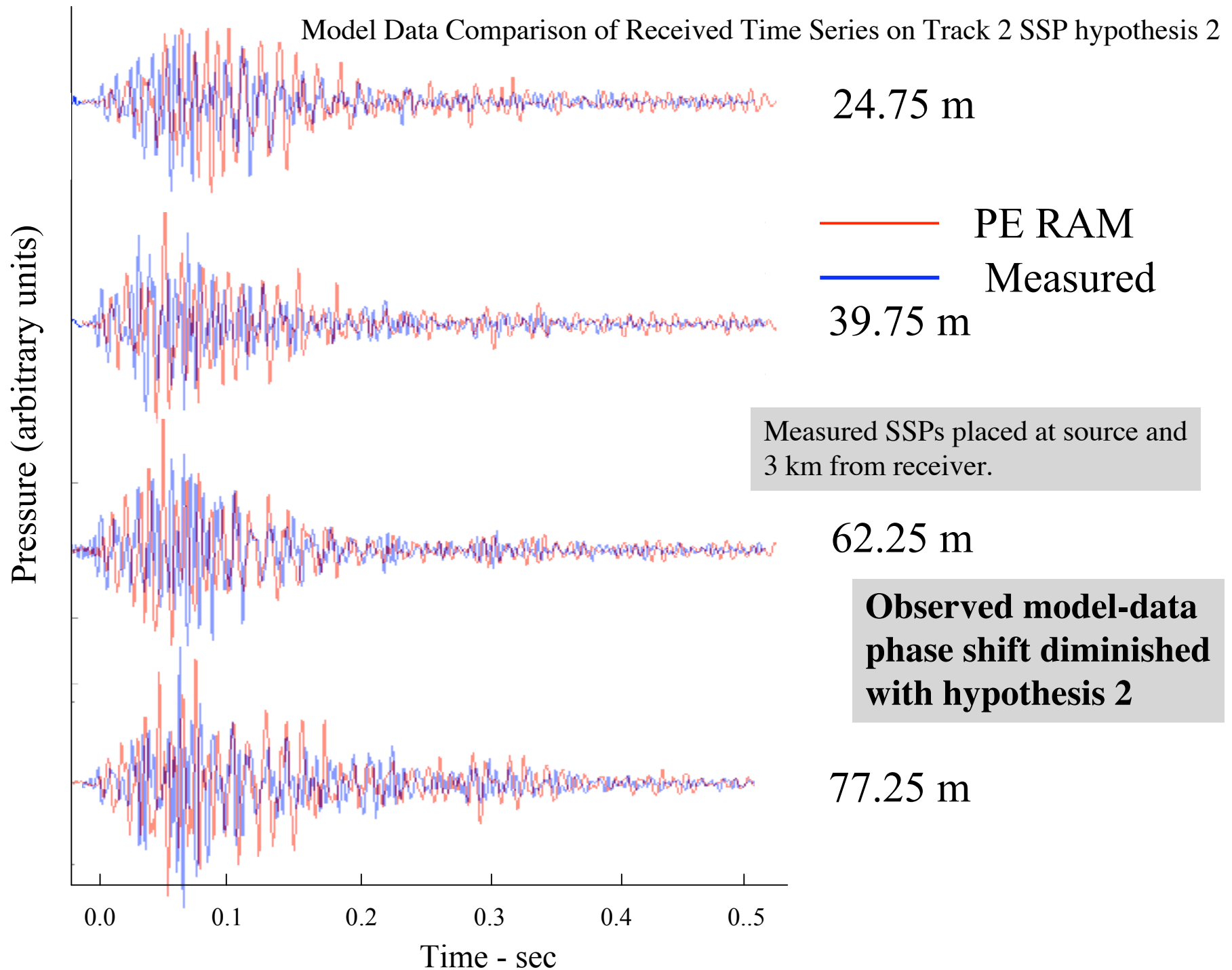
Moderately range-dependent track



Model Data Comparison of Received Time Series on Track 2 SSP hypothesis 1



Model Data Comparison of Received Time Series on Track 2 SSP hypothesis 2



Summary and Implications

- Learning probability distribution from measured acoustic data, while not the physics problem, is the technical problem
- Continued study of connections between Bayesian and Maximum Entropy Principle approach
- Details of coherent time structure of received time series for cross-shelf propagation sensitive to both the range dependence of the
 - Geoacoustic profile
 - SSP profile
- Resolving ambiguities
 - Include range-dependence of SSP and geoacoustics in $\rho(W|D)$
 - Balance of information gain and loss from increased number of parameters
 - Sampling of water column in long range LF shelf experiments needs to be not greater than 5 km
 - Implications for 50x20 km² future shelf experiment

Acknowledgements: (1) WHOI for making Array 3 data and CDT data available