

Observations and modeling of angular compression and spatial coherence in sea surface forward scattering

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Spatial coherence in forward scattering from single (time resolved)
interaction with sea surface from *Shallow Water 06*

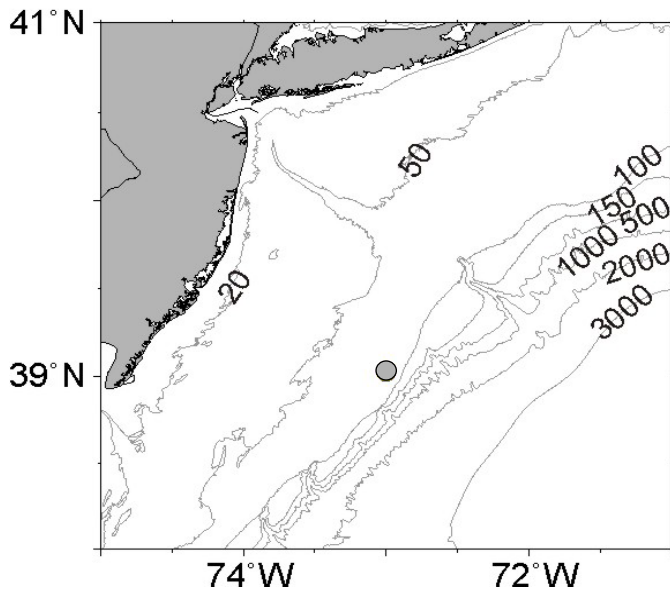
Environment: Wind speed ~ 6 m/s, Waveheight ~ 0.15 m, stationary > 6 h

Comparative influence of sea surface $C(Z)$ [thermocline]



Research sponsored by U.S. Office of Naval Research



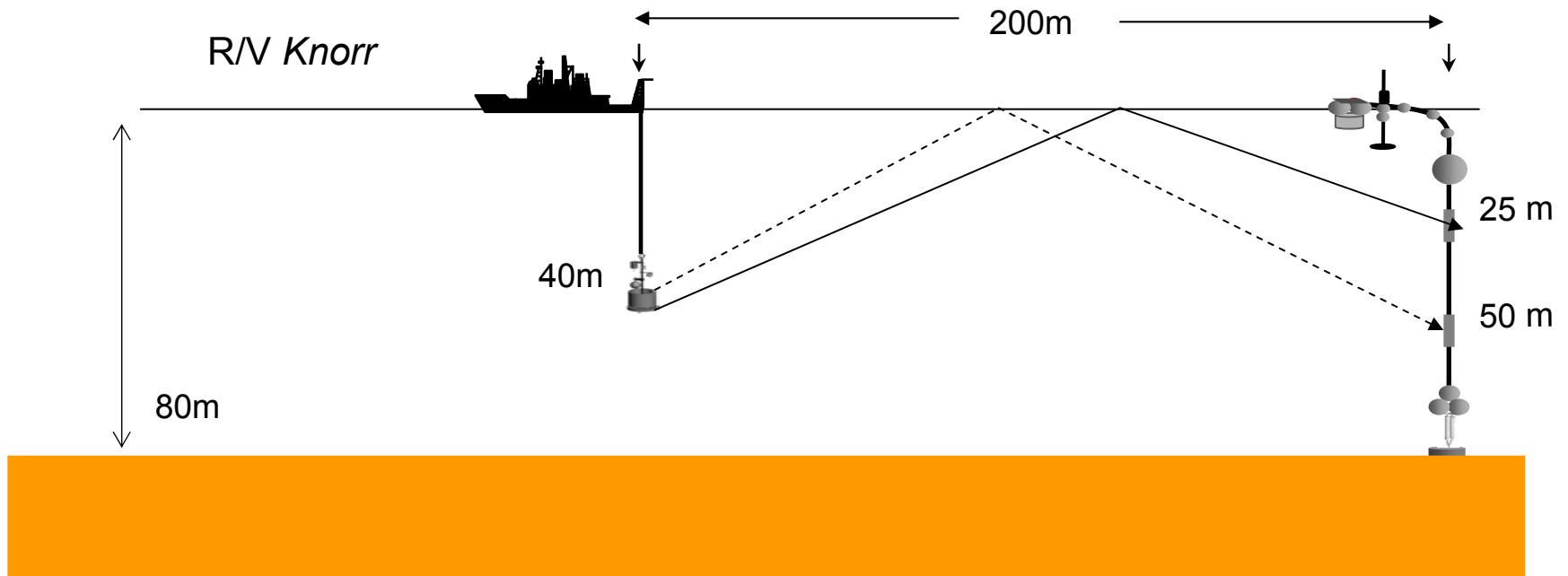


Experimental site: off the New Jersey
Continental Shelf, Water Depth 80 m

Shallow Water 06 (**SW06**) August 2006

Acoustic Source

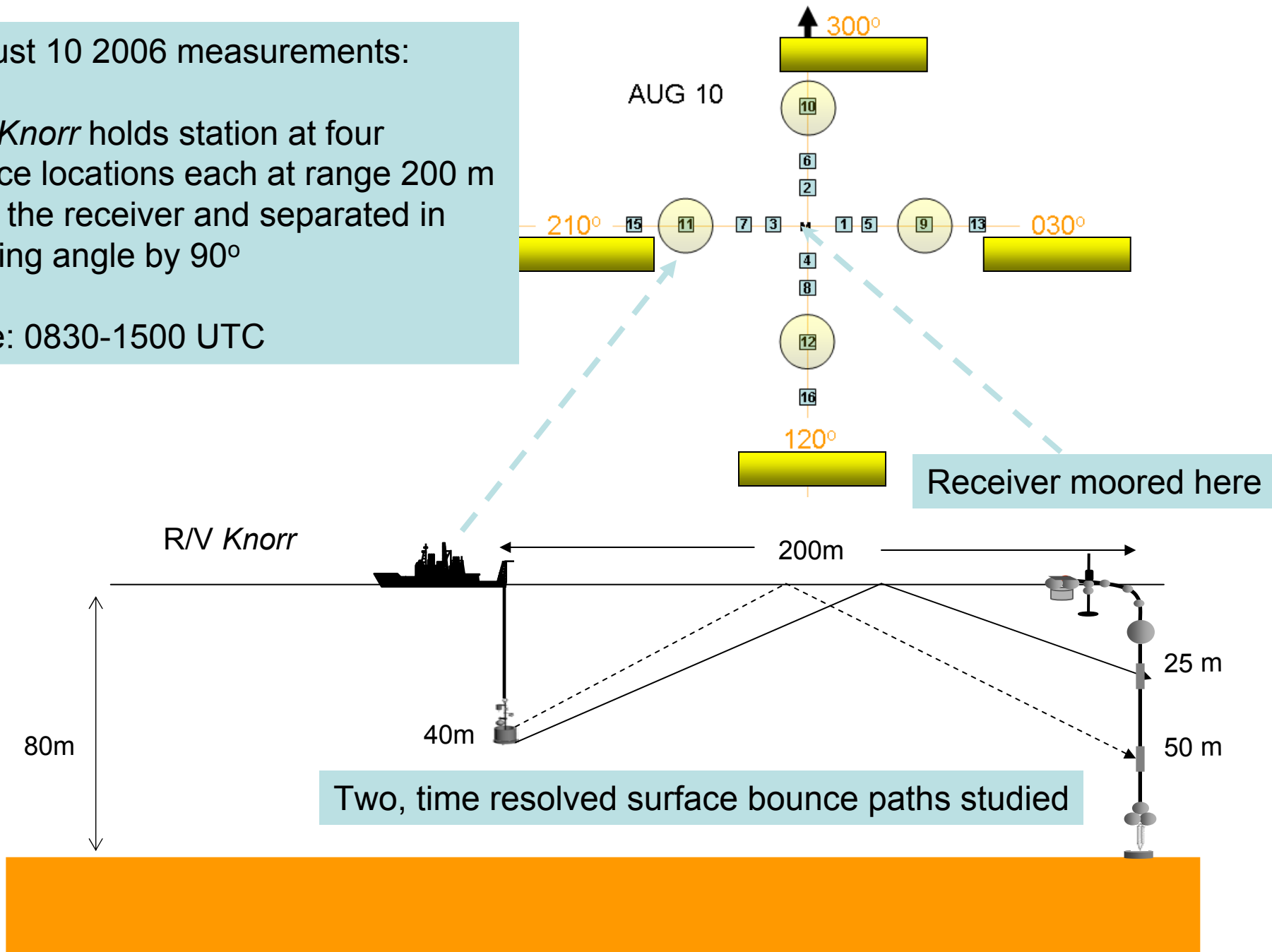
Moored Receiver
& Data Telemetry

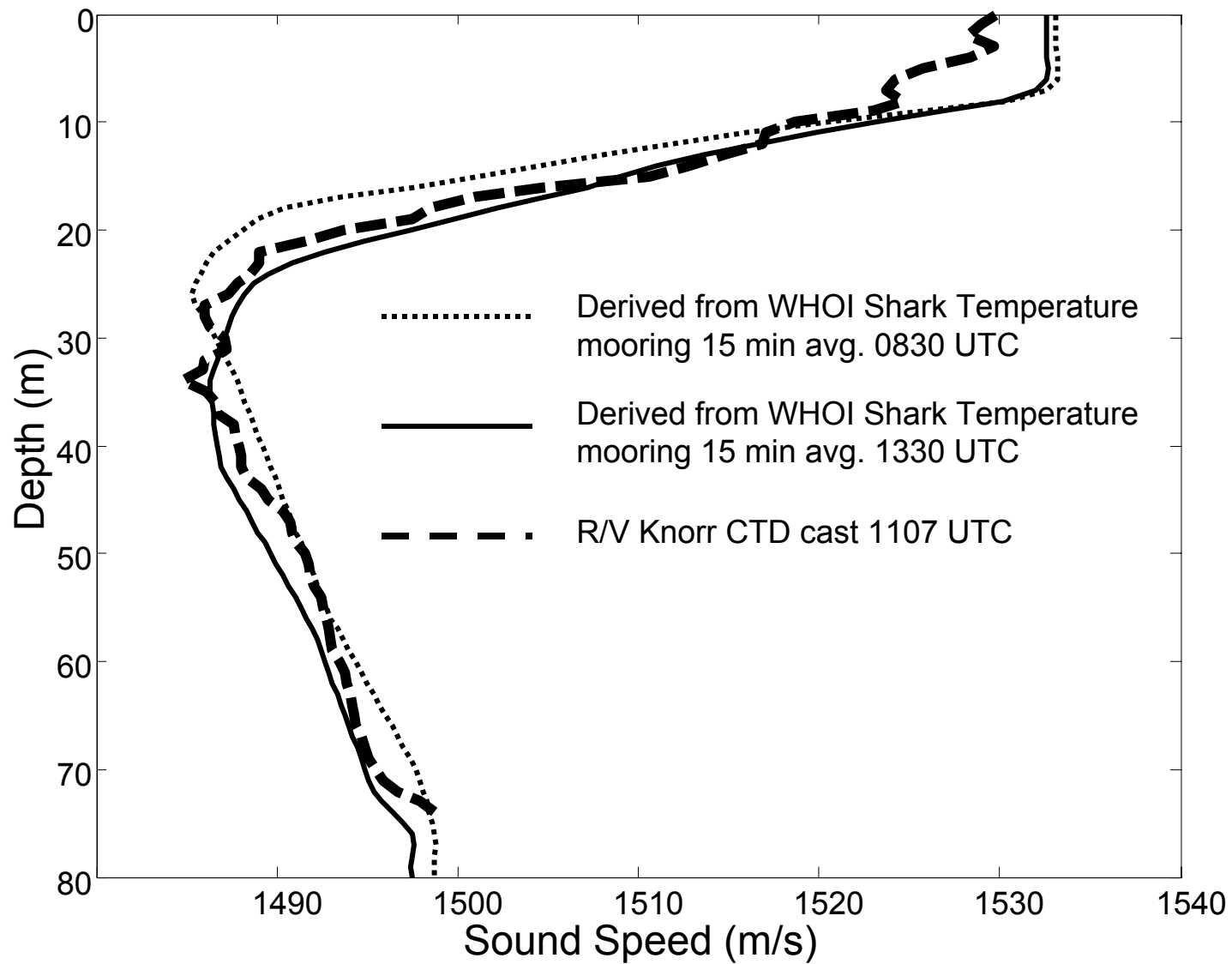


August 10 2006 measurements:

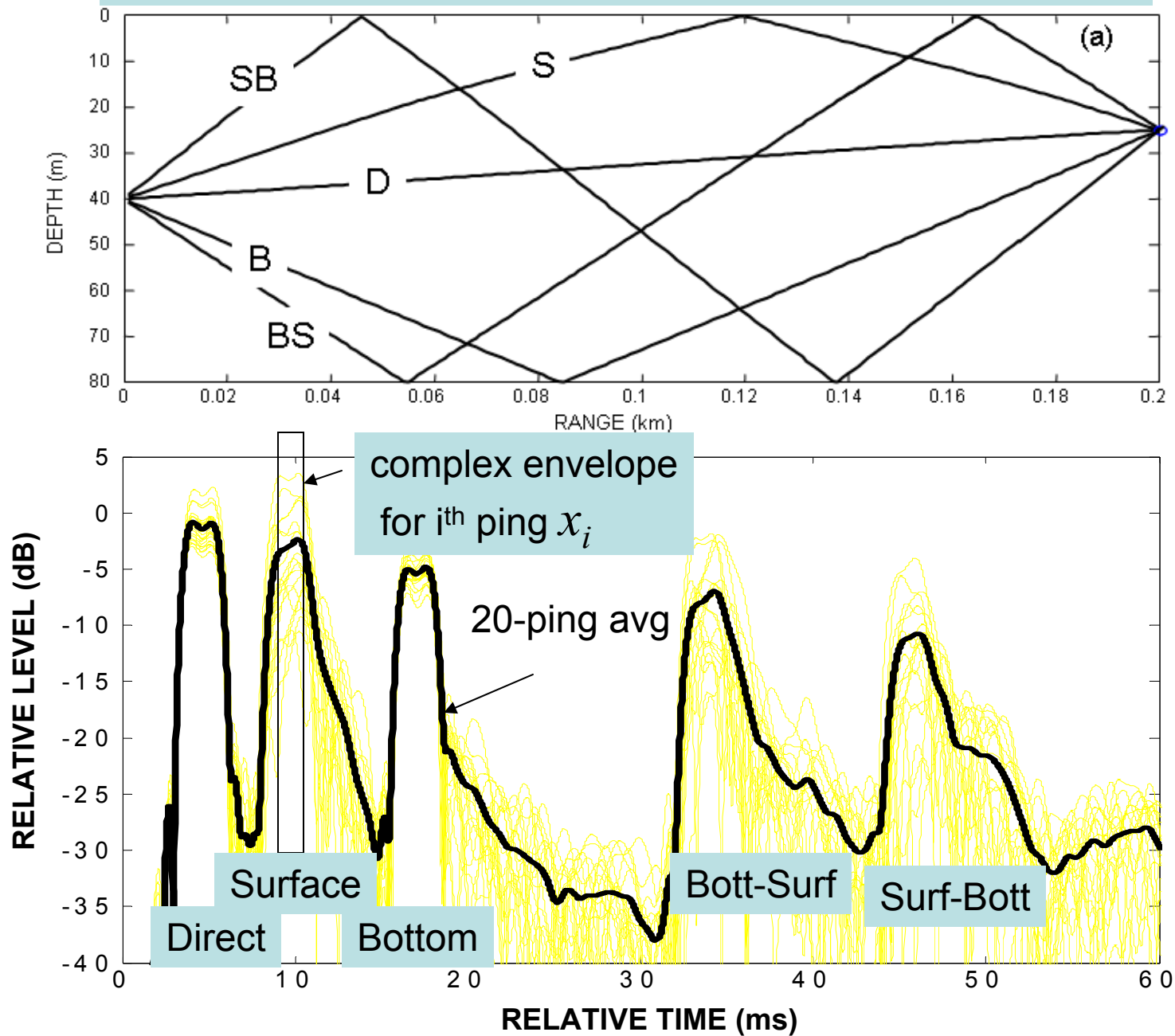
R/V *Knorr* holds station at four source locations each at range 200 m from the receiver and separated in bearing angle by 90°

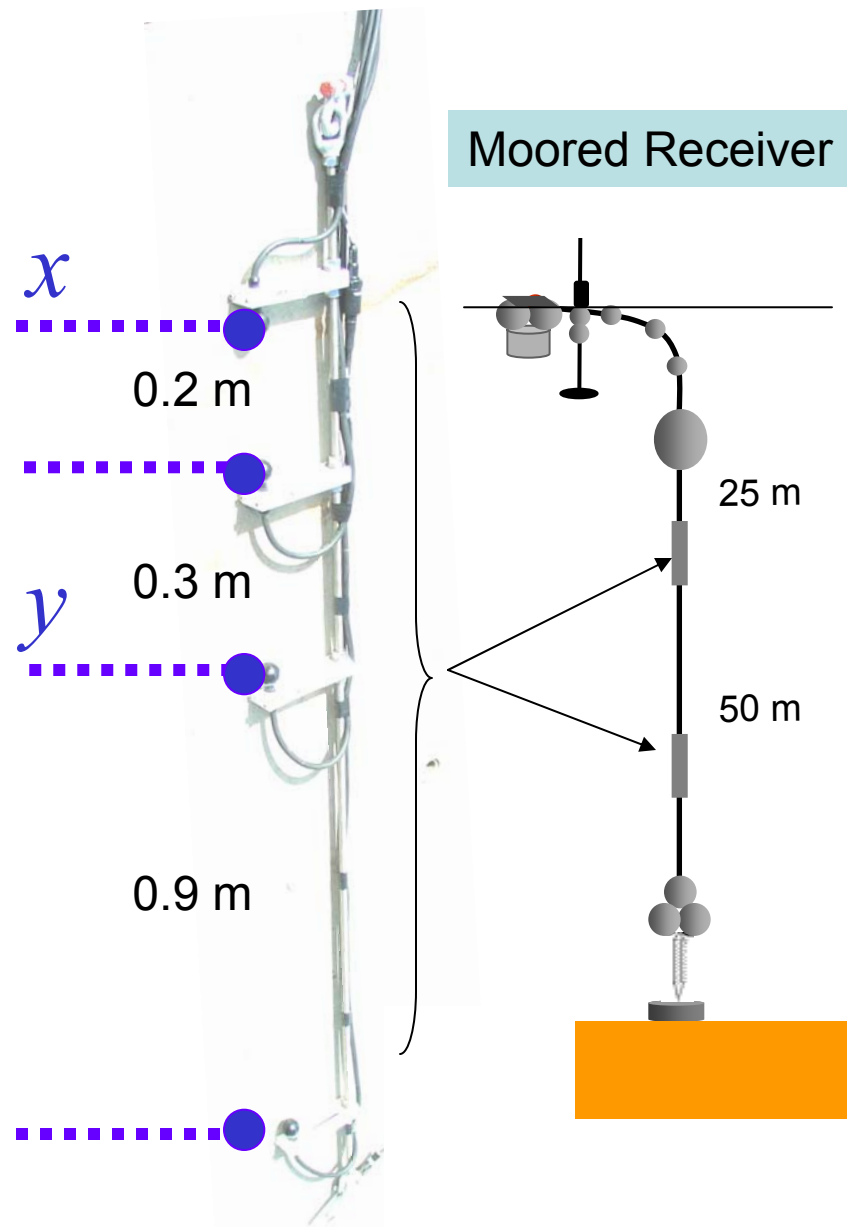
Time: 0830-1500 UTC





upper receiver eigenrays and corresponding arrival structure



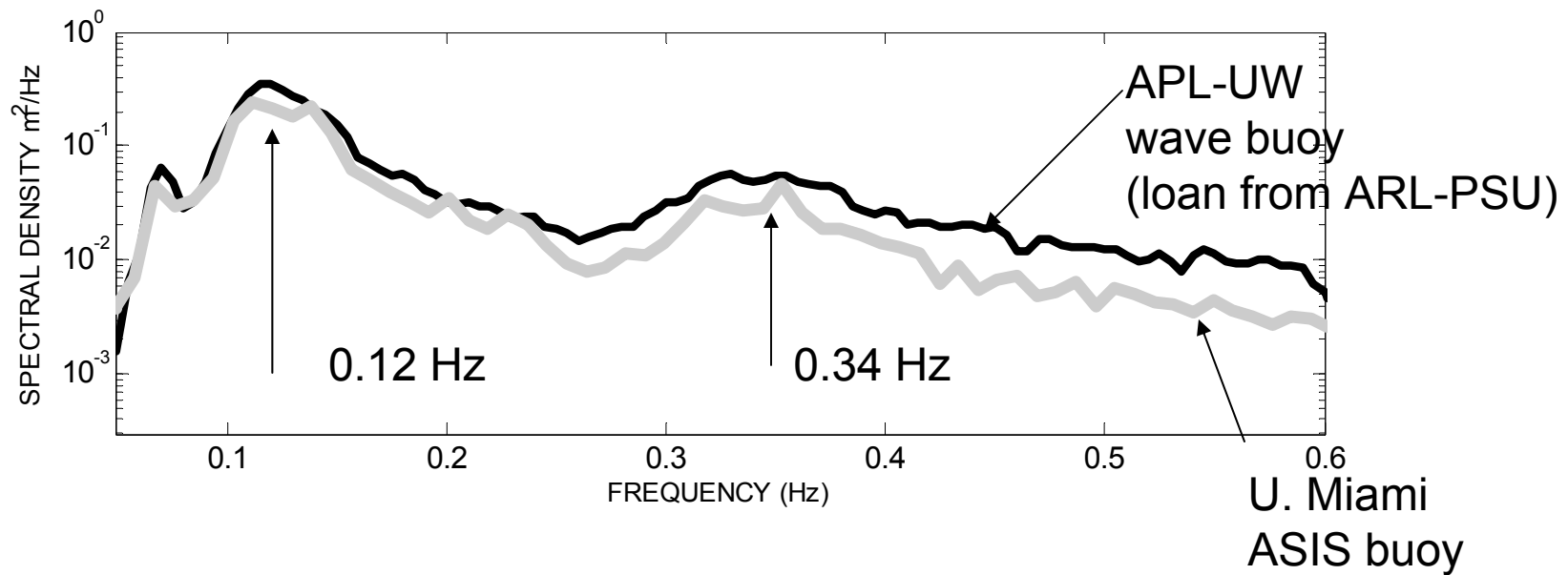
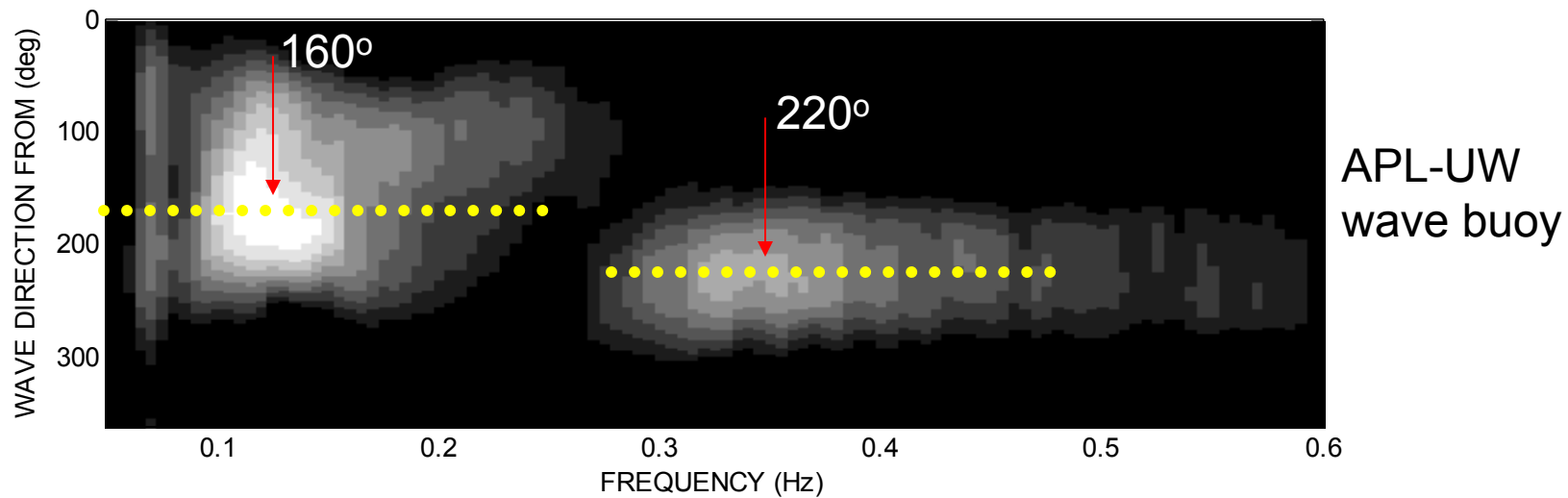


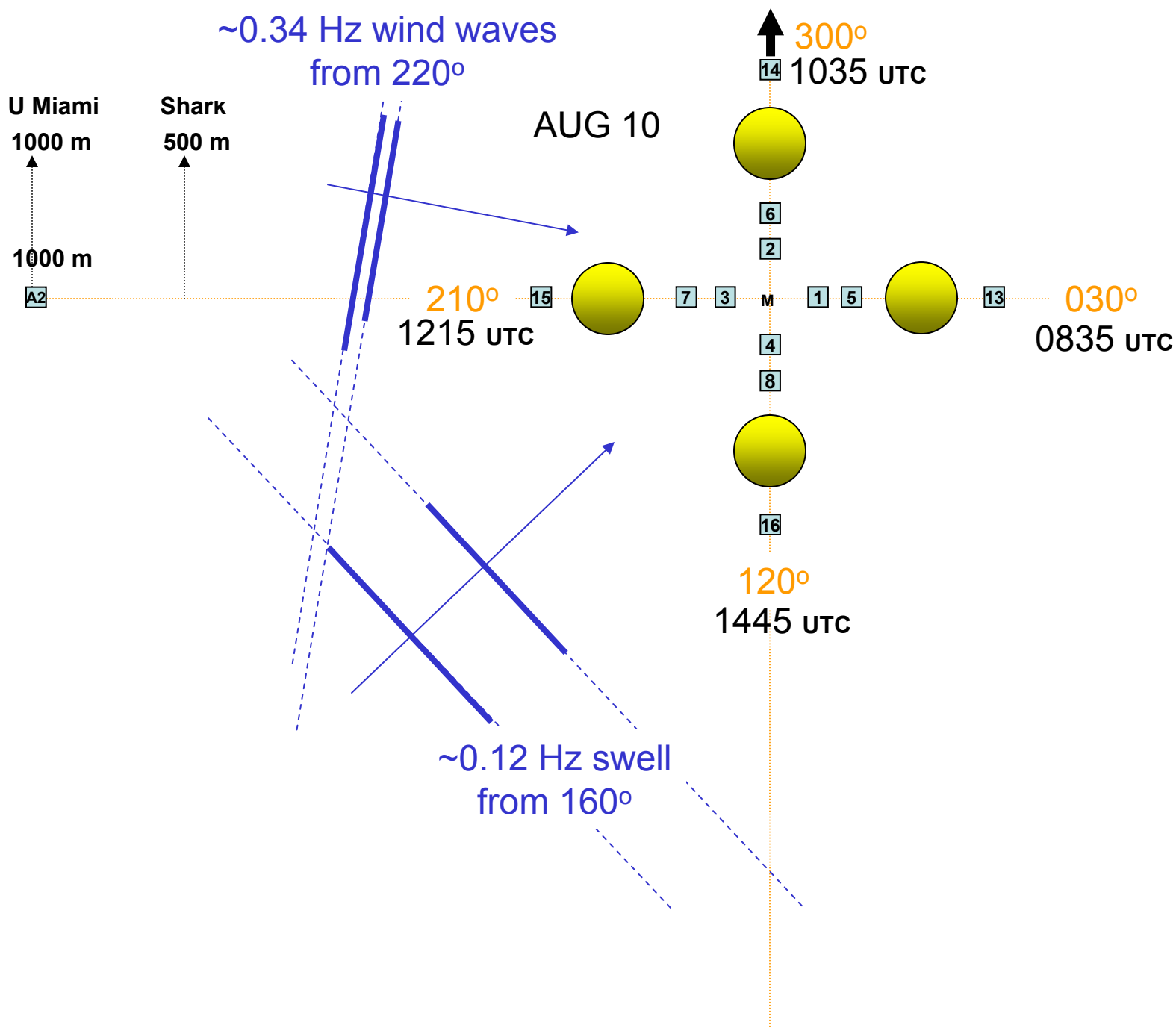
Spatial coherence between
(d) vertically-separated
channels based on 20 ping avg

$$\Gamma_{xy} = \frac{\langle xy^* \rangle}{\sqrt{\langle xx^* \rangle \langle yy^* \rangle}}$$

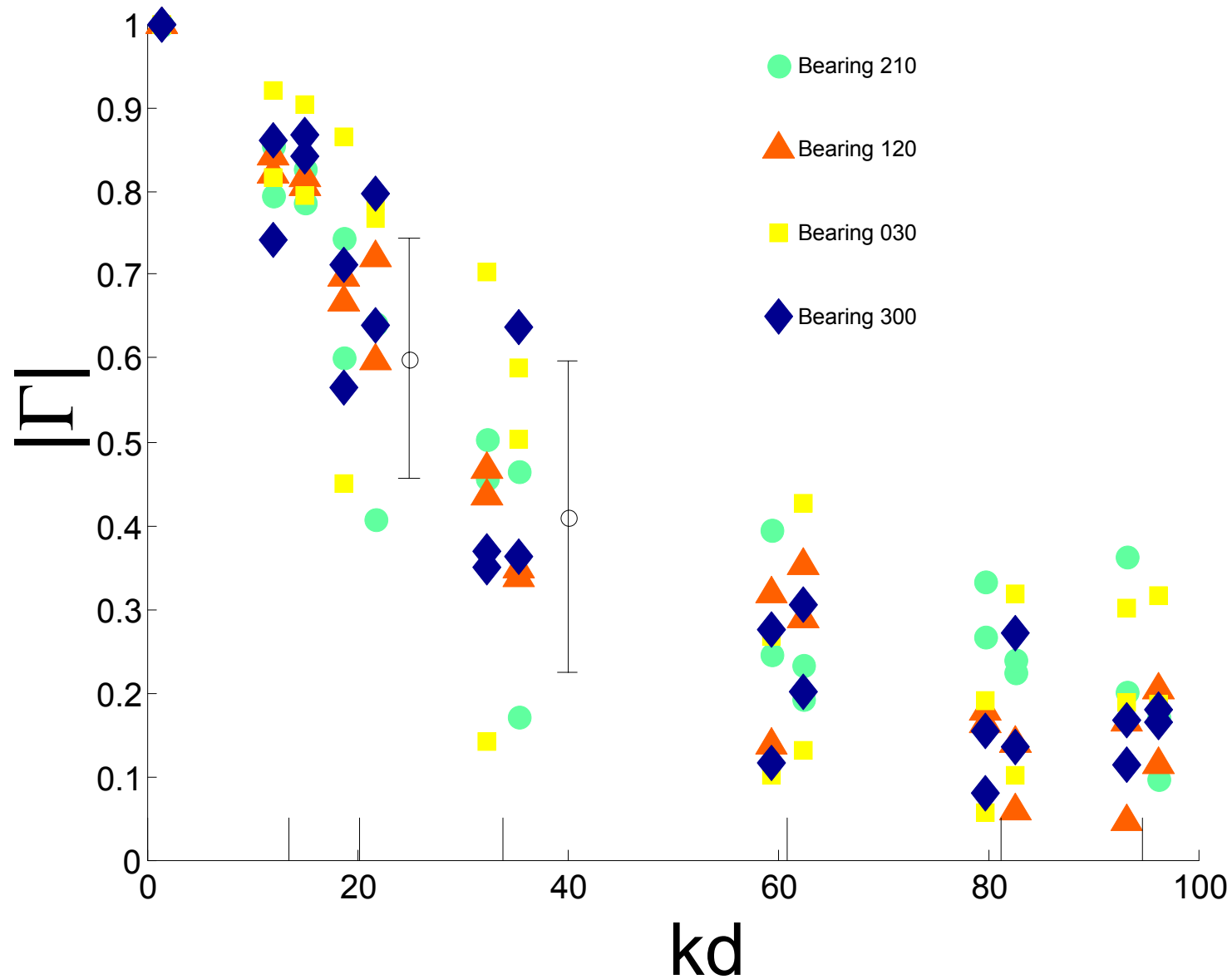
4 receiver pairs and frequency (k)
6 combinations of kd

Average air-sea conditions for 0830-1500 UTC. Wind speed 6 m/s \pm 1 m/s

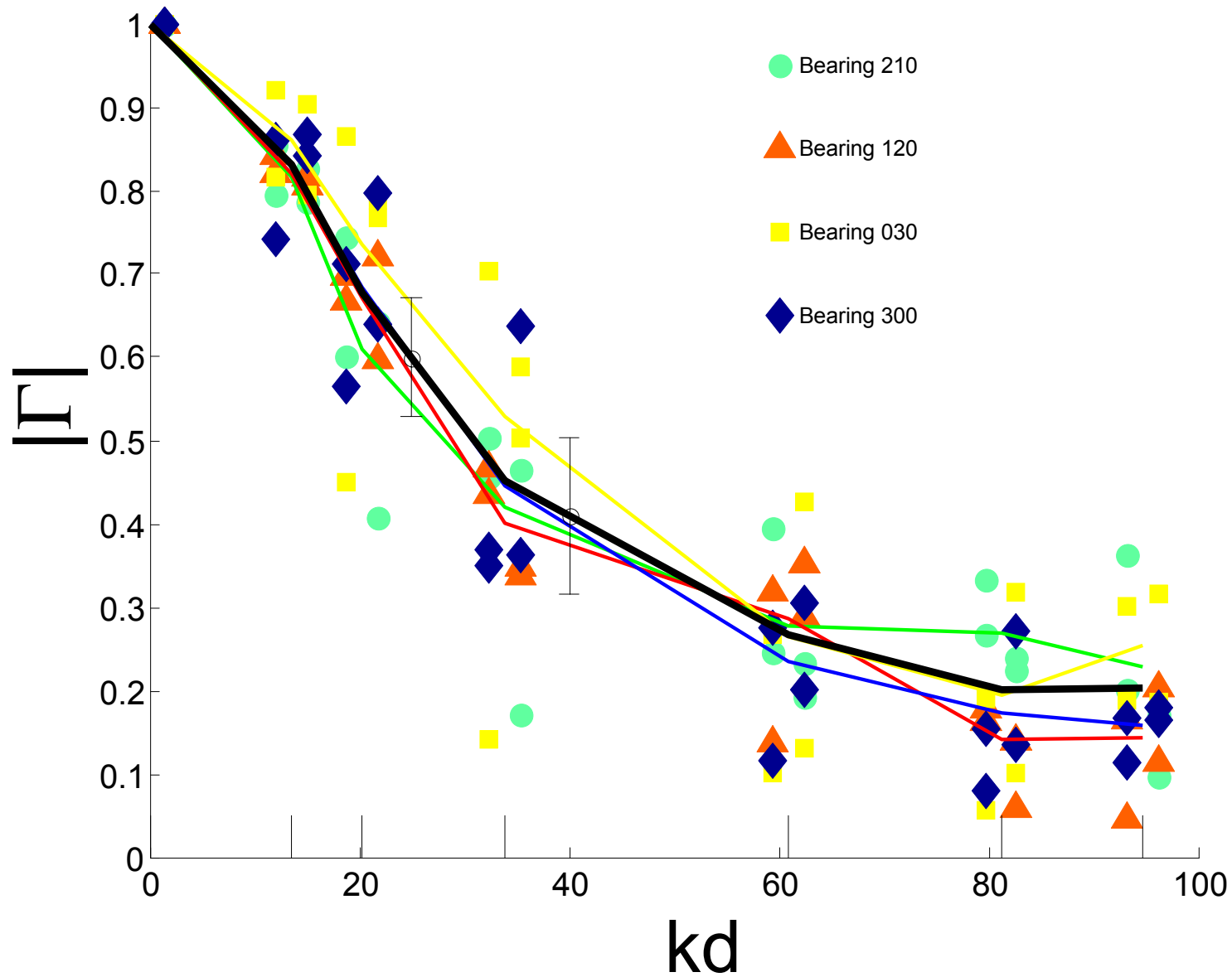




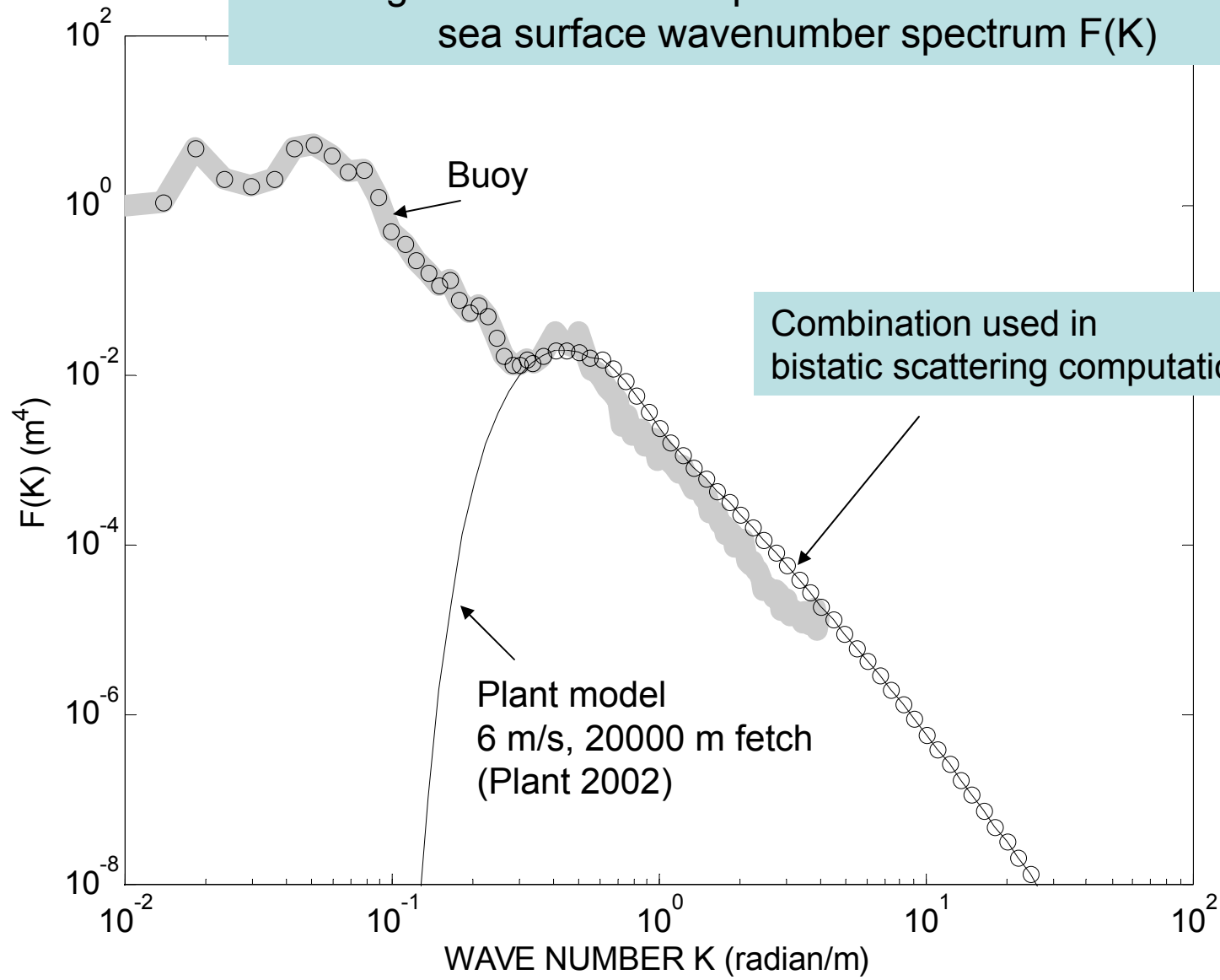
Absolute value of vertical coherence vs normalized separation (kd) at 16 kHz



Absolute value of vertical coherence vs normalized separation (kd) at 16 kHz

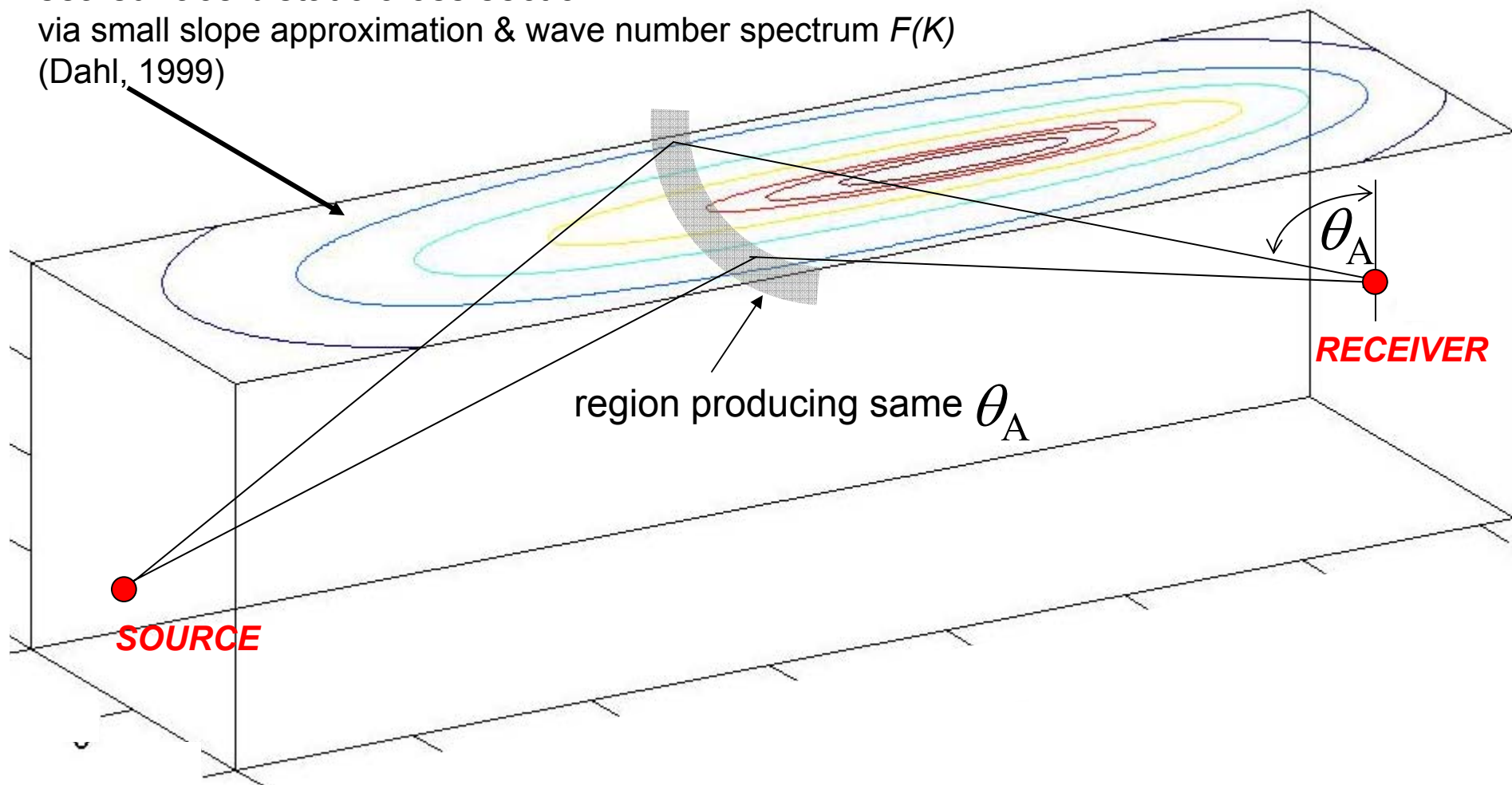


Modeling of coherence will proceed with directional-averaged
sea surface wavenumber spectrum $F(K)$

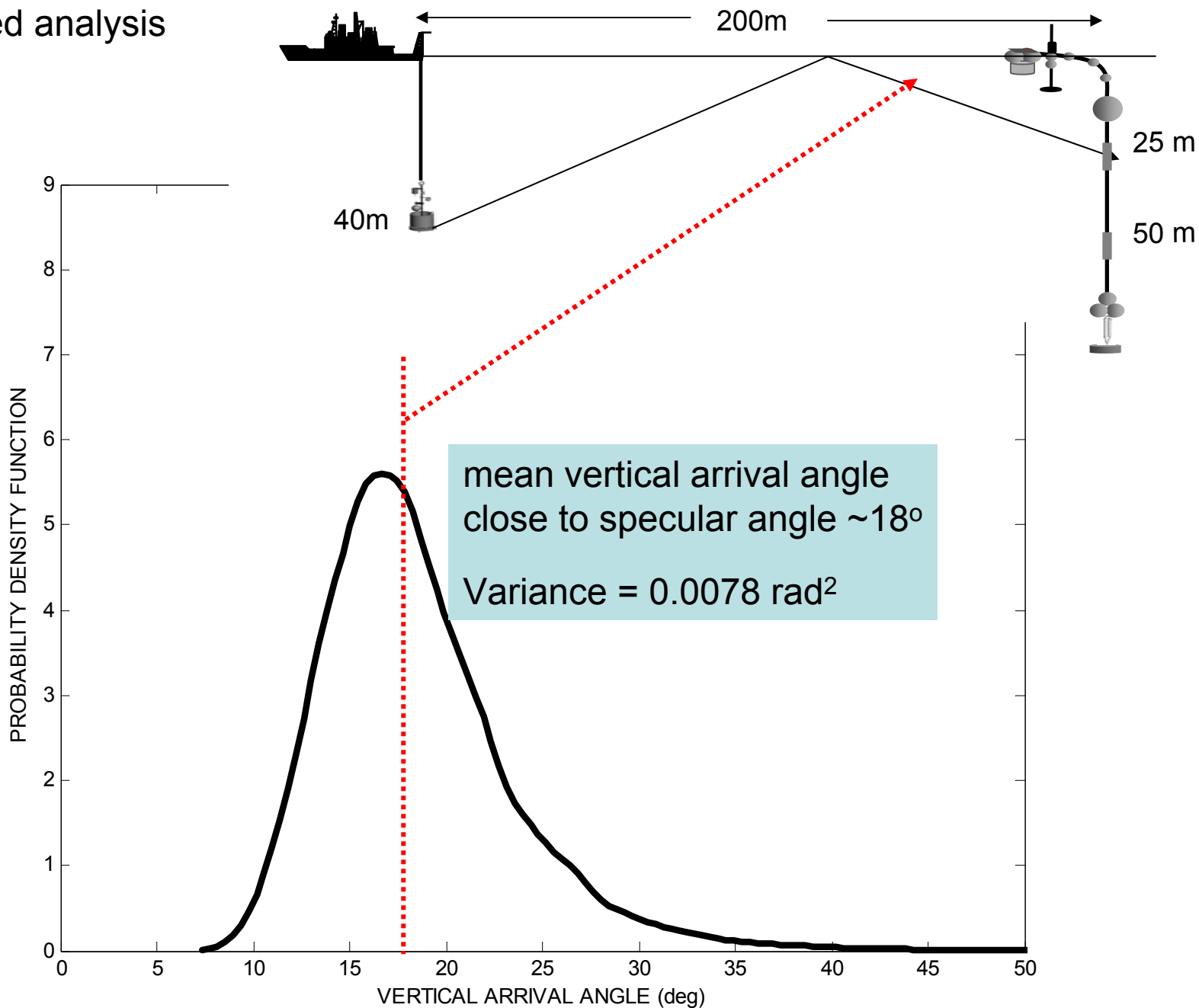


PDF for vertical arrival angle

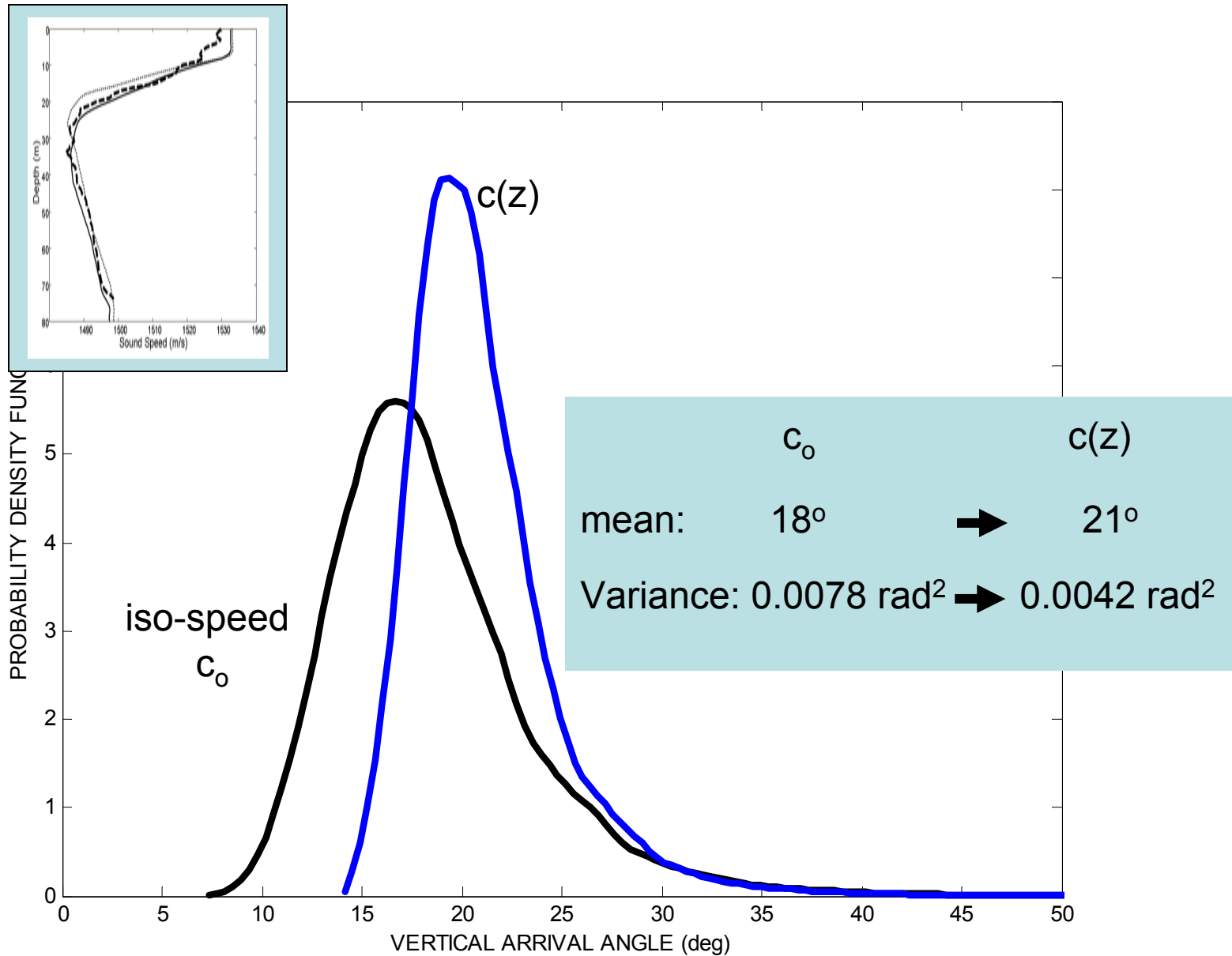
sea surface bistatic cross section
via small slope approximation & wave number spectrum $F(K)$
(Dahl, 1999)



iso-speed analysis

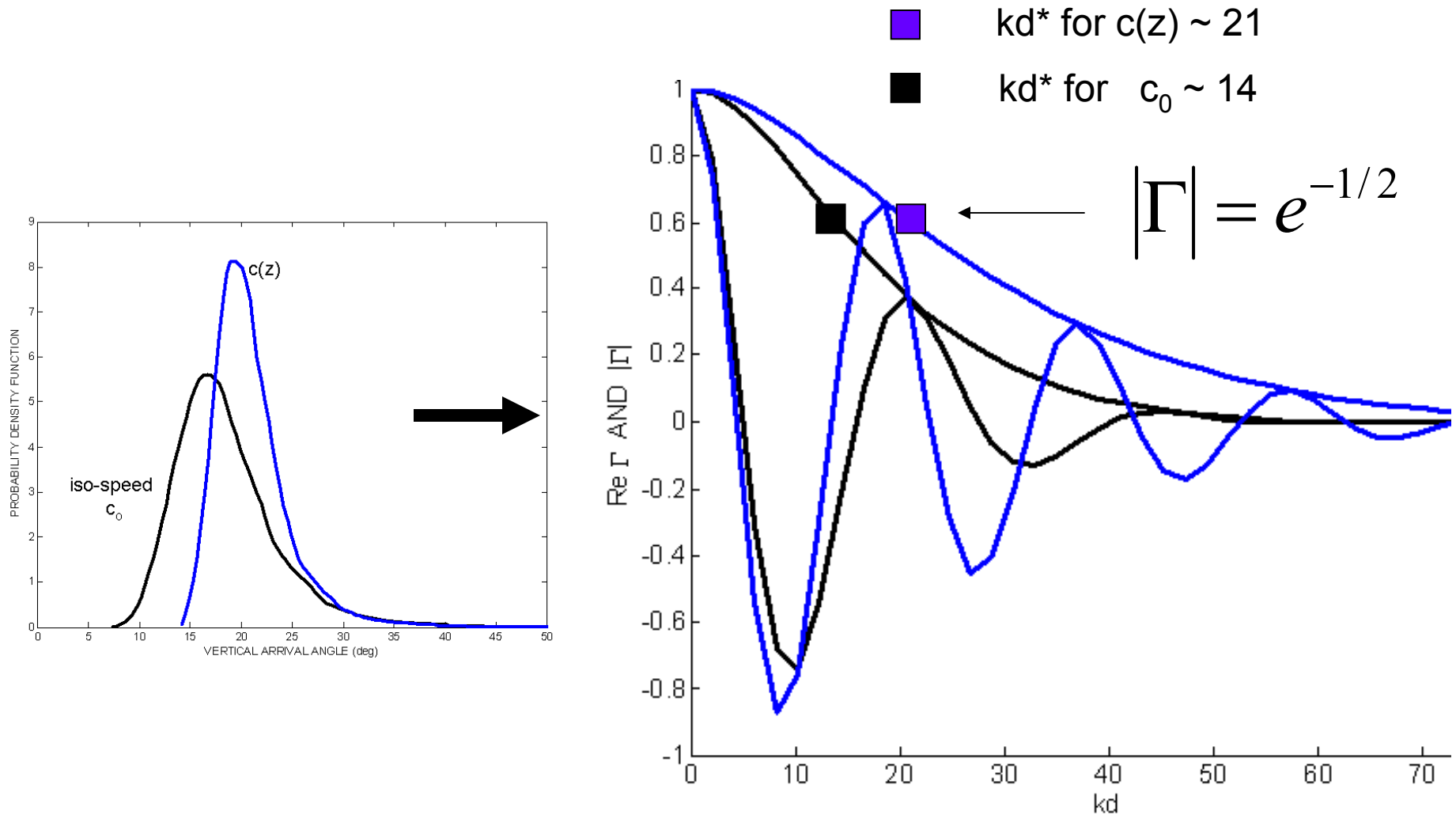


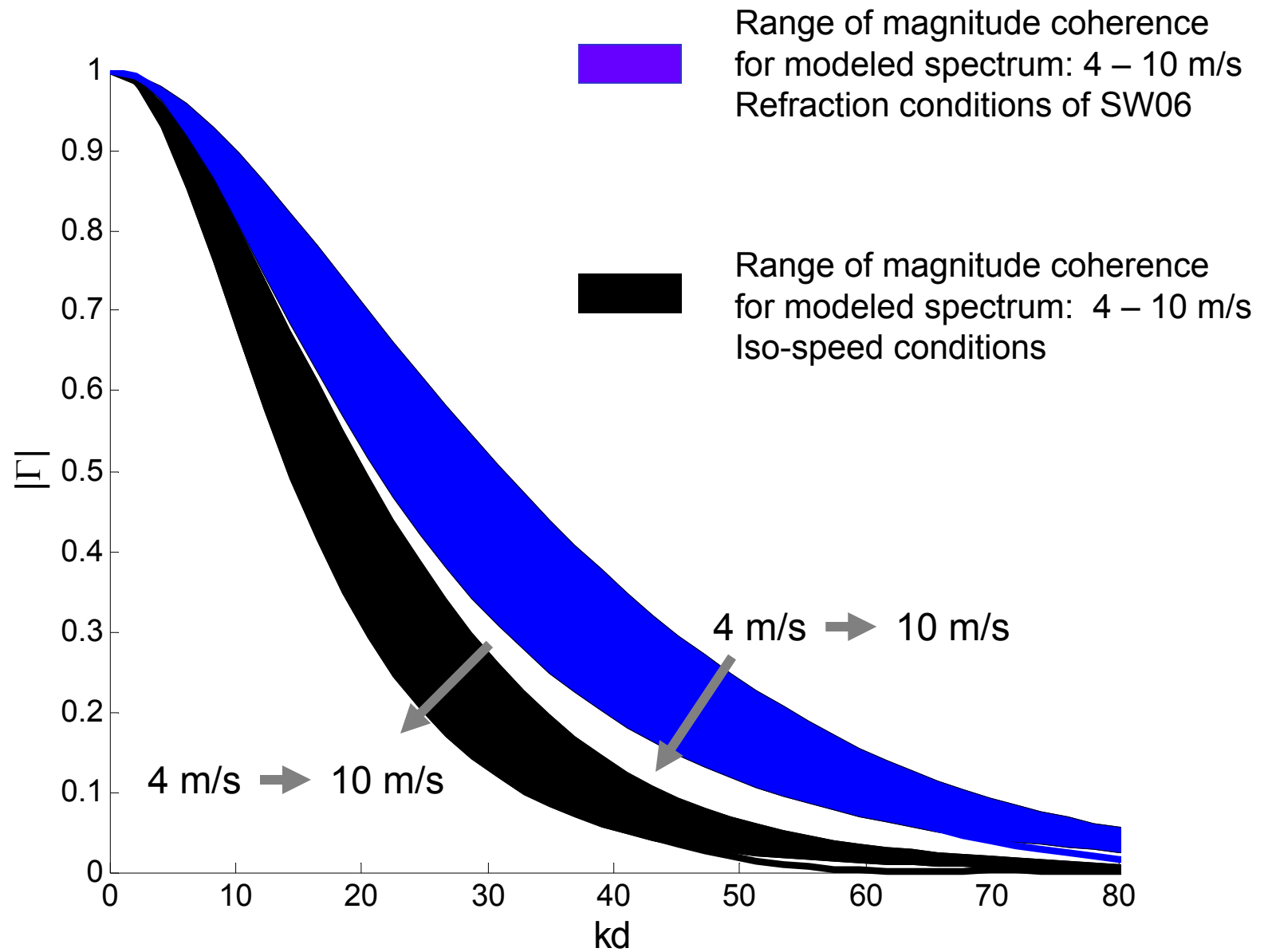
Analysis using measured $c(z)$ with thermocline

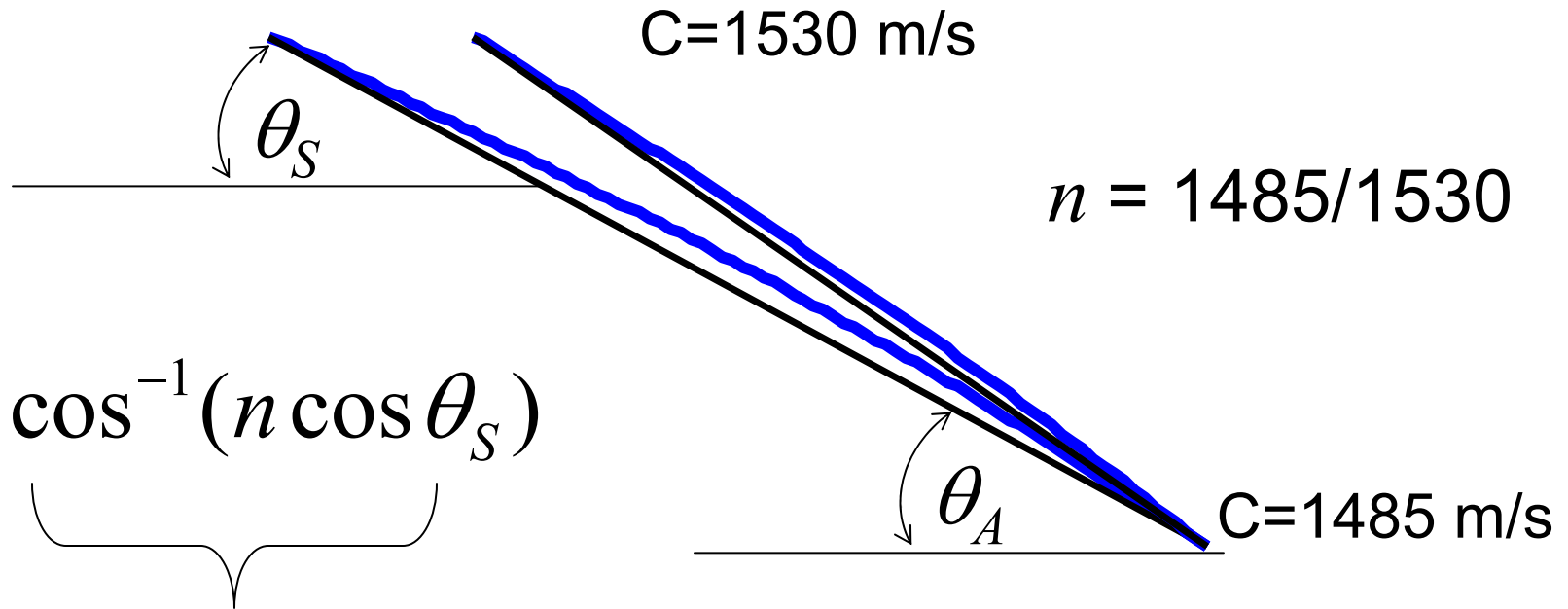


The PDF for vertical arrival angle is readily converted to spatial coherence $\Gamma(kd)$

Alternatively, the van Cittert-Zernike Theorem can be utilized to estimate $\Gamma(kd)$
(Dahl 2002, 2004)







$$\theta_A = \cos^{-1}(n \cos \theta_S)$$

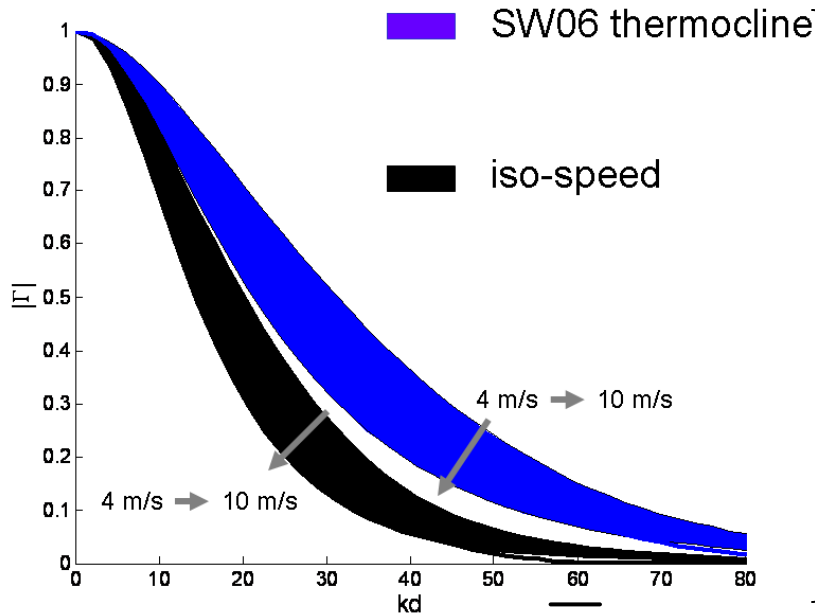
f is a smooth function relating surface-to-arrival angle

$$\therefore \text{var}(\theta_A) \approx \left(\frac{\partial f}{\partial \theta_S} \Big|_{\bar{\theta}_S} \right)^2 \text{var}(\theta_S)$$

Vertical angular compression factor

$$\frac{-n \sin(\bar{\theta}_S)}{\sqrt{1 - n^2 \cos^2 \bar{\theta}_S}}$$

Vertical Angular Compression



Large change in kd^* predicted by the angular compression factor

Compression does *not* \Rightarrow \uparrow intensity

SW06 geometry:
TL *increased* by 1.5 dB (confirmed by ray and PE analysis)

$$\sigma_{\theta_A} = \frac{n \sin(\bar{\theta}_S)}{\sqrt{1 - n^2 \cos^2 \bar{\theta}_S}}$$

SW06 thermocline

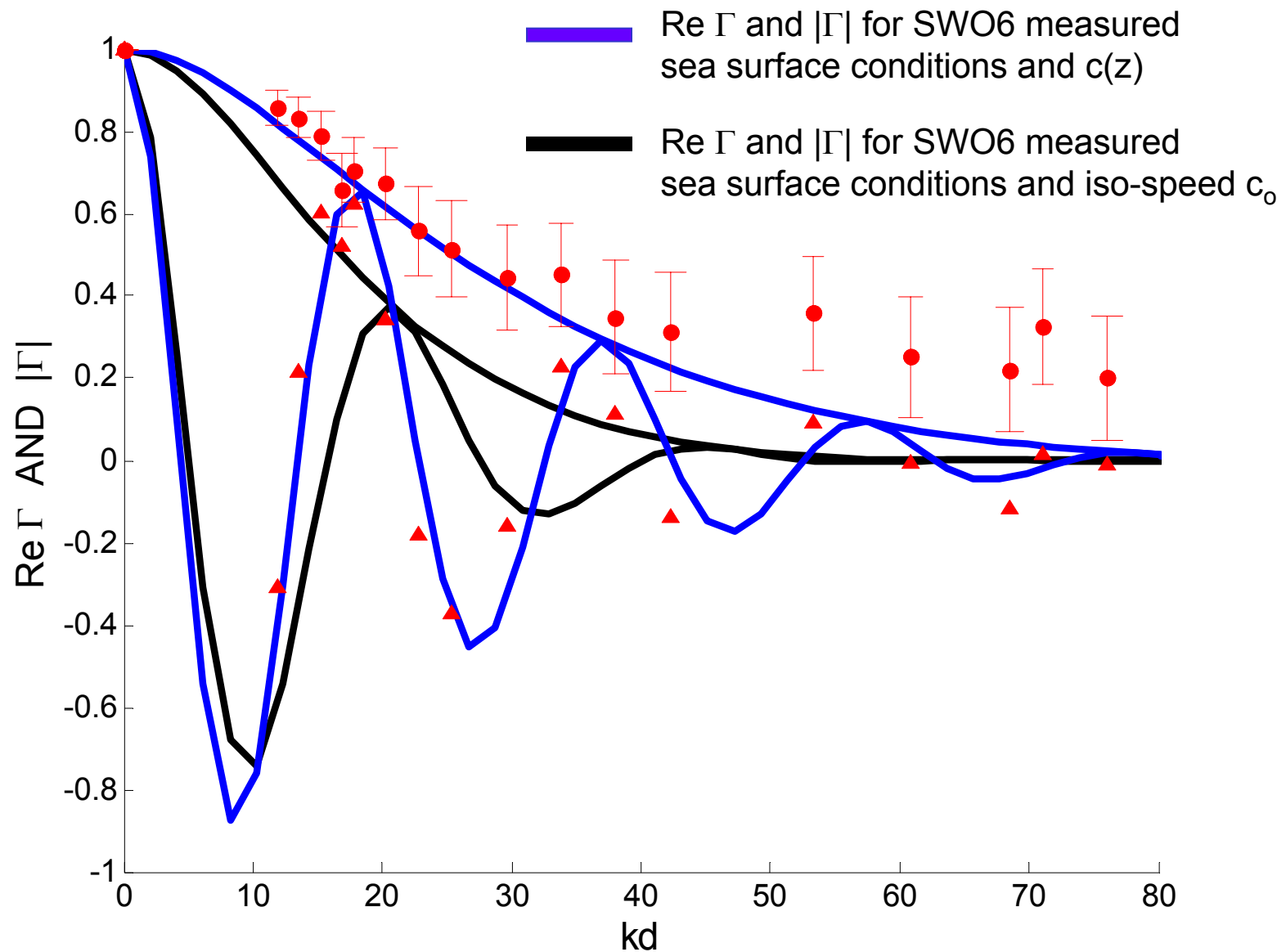
0.72

σ_{θ_A}
iso-speed

$$kd^* \sim 1 / \sigma_{\theta_A}$$

$\therefore kd^*$ should \uparrow by $\sim 4/3$

Model comparison with data (14-16-18-20 kHz) plotted versus kd



Summary

- Spatial coherence in sea surface forward scattering with strong thermocline
- Vertical angular compression: dominate effect greater than that linked to sea surface roughness and slope
- Vertical angles *compressed* while TL *increased* over spherical spreading (angle expansion in upward paths not balanced by downward paths)
- Mild refraction effects (influencing phase of Γ) observed in ASIAEX data (Dahl 2004)
SWO6: strong refraction effects influencing both magnitude and phase
- Predictive model based on Snell's mapping of angular variances

